This manuscript version is made available under the CC-BY-NC-ND 4.0 license http://creativecommons.org/licenses/by-nc-nd/4.0/

Revised manuscript submitted to Chemosphere 1 2 2 September 2015 3 Colloidal mobilization of arsenic from mining-affected soils by 4 surface runoff 5 6 7 Miguel Angel Gomez-Gonzalez^a, Andreas Voegelin^b, Javier Garcia-Guinea^a, Eduardo 8 Bolea ^c, Francisco Laborda ^c, Fernando Garrido ^{a*} 9 10 11 12 ^a Museo Nacional de Ciencias Naturales (MNCN, CSIC).C/JoseGutierrez Abascal 2, 13 14 28006, Madrid, Spain ^b EAWAG, Swiss Federal Institute of Aquatic Science and Technology. 15 Ueberlandstrasse 133, 8600, Duebendorf, Switzerland 16 ^cInstituto Universitario de Ciencias Ambientales (IUCA), Universidad de Zaragoza. C/ 17 Pedro Cerbuna 12, 50009, Zaragoza, Spain 18 19 20 Keywords: Arsenic; runoff; colloids; AF4-ICP-MS; XAS; Rainfall simulation 21 22 *corresponding author.fernando.garrido@mncn.csic.es 23

ABSTRACT

Scorodite-rich wastes left as a legacy of mining and smelting operations pose a threat to environmental health. Colloids formed by the weathering of processing wastes may control the release of arsenic (As) into surface waters. At a former mine site in Madrid (Spain), we investigated the mobilization of colloidal As by surface runoff from weathered processing wastes and from sediments in the bed of a draining creek and a downstream sedimentation-pond. Colloids mobilized by surfacerunoff during simulated rain events were characterized for their composition, structure and mode of As uptake using asymmetricflow field-flow fractionation coupled to inductively plasma mass spectrometry (AF4-ICP-MS) and X-ray absorption spectroscopy (XAS) at the As and Fe K-edges. Colloidal scorodite mobilized in surface runoff from the waste pile is acting as a mobile As carrier. In surface runoff from the river bed and the sedimentation pond, ferrihydrite was identified as the dominantAs-bearing colloidal phase. The results from this study suggest that mobilization of As-bearing colloids by surface runoff may play an important role in the dispersion of Asfrom metallurgical wastes deposited above ground and needs to be considered in risk assessment.

1. INTRODUCTION

Mining and smelting operations have changed the global distribution and occurrence of metal (oid)s at the Earth's surface (Rauch, 2012), and the release of metals and metalloids from mining and processing wastes into aquatic and terrestrial environments may threaten environmental and human health. Human activities have led to substantial As contamination in the environment (Morin and Calas, 2006), as often occurs in association with economically relevant metal ores. Accordingly, the roasting and

smelting of metal ores has led to widespread environmental contamination via 50 atmospheric emissions and deposition of waste materials(Vaughan, 2006). In Spain, 51 there are many As-affected legacysites affected from former miningactivities. Often, 52 53 As-bearing processing wastes had beendumped aboveground and remain exposed to rainfall and weathering, posing a risk for adjacent water bodies(Garcia-Sanchez and 54 Alvarez-Ayuso, 2003). 55 56 The weathering of processing sulfide-containing wastes leads to the formation of acid 57 mine drainage (AMD). The oxidation of Fe(II) and precipitation of Fe(III) in neutralized AMD may lead to the formation of mobile colloidal carriers for As (Cheng et al., 58 2009). Due to their physicochemical properties, colloids may persist in aqueous 59 suspensions over extended periods of time andact as carriers for poorly soluble metals 60 and metalloids in surface and surbsurface environments(Kretzschmar and Sticher, 61 62 1997). The affinity of As to colloidalFe(III) and the importance of such colloidal particles as As-mobilizing carriers in the environment has been documented(Fritzsche et 63 64 al., 2011). For instance, Grosbois et al.(2011)identified different As-carriers during the transport of solid particulate matter in varying abundances according to the hydrological 65 cycle of a river draining a former gold mine district in France. Furthermore, colloids and 66 colloid-associated contaminants, mobilized by rainfall within ephemeral overland 67 courses, can be transported rapidly through the vadose zone with minimal interaction 68 with the soil matrix (Ranville et al., 2005). This can be attributed to the presence of 69 preferential flow paths, which in turn may play a role in the solid-phase As distribution 70 71 and Asretention mechanism in soils, especially in locations affected by large amounts of metals released with AMD(Helmhart et al., 2012). Overall, the importance of surface 72 73 runoff as compared to infiltration for As dispersion may vary with rainfall intensity and duration, vegetation cover, soil topography, as well as soil physical properties and flow 74

conditions through soils and riverbeds. Irrespective of the mode of colloid transport, the 75 nature and stability of As-bearing colloids in tailings and mining wastes that control the 76 extent of As release needs to be assessed. 77 78 The combined use of spectroscopic, spectrometric and fractionation techniques offers the means to gain information on the size-dependent elemental composition of colloids, 79 the nature of the colloidal carrier phase, and the speciation of associated contaminants. 80 Rainfall simulation experiments on the other hand, allow to control rainfall parameters 81 (intensity and duration) and to exclude the inherent variability associated with natural 82 rain that increases the complexity of stormwater quality research (Bian et al., 83 84 2011). Simulated rainfall have been used to assess heavy metal contamination in urban areas(Wicke et al., 2012), but only few studies have been performed on contaminated 85 soils under natural conditions(Fernandez-Galvez et al., 2008). To date, 86 87 however, the colloid-facilitated release of Aswith surface runoff in mining-affected areas has not been assessed. 88 Asymmetric flow field-flow fractionation (AF4) is a separation technique for the 89 90 characterization of colloids from natural samples(Laborda et al., 2011). When coupled to an inductively coupled plasma mass spectrometer (ICP-MS), the size-91 dependentelemental composition of the colloidal phasescan be determined. For instance, 92 Neubauer et al.(2013) characterized the colloid distribution in soil runoff as generated 93 by storm events using AF4-ICP-MS. The authors showed that a colloidal fraction 94 ranging from 0.2 to 0.45 µm was mobilized but they did not assess the speciation of the 95 96 colloidal As. Direct insight into the speciation of colloidal As and Fe can be gained using X-ray absorption spectroscopy (XAS), widely used to study the speciation of As 97 and Fe in natural samples(O'Day et al., 2004; Voegelin et al., 2007),includingdispersible 98 soil colloids(Regelink et al., 2014; Serrano et al., 2015). 99

In previous studies, we examined the dispersion and solid phase partitioning of As in soils and sediments affected by weathering and erosion of metallurgical processing wastes at a former mining and smelting site in Madrid province (Gomez-Gonzalez et al., 2014). In the same area, the molecular-scale speciation of As associated to colloids detached from soil extractshas previously been studied (Serrano et al., 2015). In this continuing work, we aimed at determining the extent and mode of colloidal As mobilization with surface runoff during simulated rain events. AF4-ICP-MS was used to determine the colloid size distribution and As partitioning between the dissolved and colloidal phase, and Fe and As K-edge XAS to determine the colloidal carrier phase and colloidal As speciation.

2. MATERIALS AND METHODS

2.1. Site description

The experiment was conducted in a sub-catchment of the Guadalix River (Madrid, Spain) at the foot of the 'Sistema Central' mountain range(40°45'34.33" N - 3°41'07.13" O, 929 meters altitude). In the area,metal sulfides and wolframite ores in association with quartz veins were mined for tungsten extraction during the Second World War. Theabandoned smelting factory along with mining wastes deposited above ground still remain at this site (Helmhart et al., 2012). The massive pile of processing wastes (approximately 6x6 m, 1 meter-thick, 19 g As kg⁻¹), is subjected to weathering and erosion (Gomez-Gonzalez et al., 2014). The site is representative for many other legacy sites in Madrid province and other regions of Spain.

The average annual rainfallin the area reaches 705 mm and the average temperature, 13.3°C(Spanish Agency of Meteorology, AEMET). About 188-201 mm of the annual rain precipitate in winter, 197-213 mm in spring, 85-93 mm in summer and 200-233 in

autumn. Rainfall simulation tests were performed at: A) the arsenic-bearing waste pile (WP); B) the river bed (RB) of a small stream (~ 1 m wide) that seasonally collects surface runoff from WP; and C) the sediment that accumulatesdownstream in an artificial sedimentation pond (SP) (Figure S1). During intense rainfall, a major part of runoff waters is usually retained in the sedimentation pond remaining stagnant over several days until evaporation and infiltration lead to the drying out of the pond. Physical, chemical, and mineralogical properties of the topmost layer of soils and sediments at the experimental locations are in Table 1. 2.2. Rainfall simulation experiments

133

134

135

136

137

138

139

125

126

127

128

129

130

131

132

Simulation experiments were performed using the portable rainfall simulator designed byCalvo et al.(1988) and modified byCerda et al.(1997). Single rainfall events of distilled water were applied at 28 mm h⁻¹ during 60 minutes in each experimental location. The basic components of the simulator and the experimental conditions are described in the Supplementary Material(FiguresS2 and S3).

140

141

142

143

144

145

146

147

148

149

2.3. Isolation of the colloid-suspensions in runoff

Three fractions of the runoff suspensions from the rainfall simulations were analyzed: (i) the colloid-containing suspension (CS, ≤ 1000 nm) obtained after centrifugation of the runoff suspension for the removal of particulate material, (ii) the solid colloids (1000-10 nm) isolated from the colloid-containing suspension by ultrafiltration using 10-nm membranes, and (iii) the dissolved fraction (DF) that passed 10-nm ultrafiltration membranes. The colloid-suspension (CS, \leq 1000 nm) thus includes both the solid colloids (1000-10 nm) and the dissolved fraction (\leq 10 nm). Runoff samples were subjected to the protocol described in the supplementary material (Figure S4).

150	
151	2.4. Size chard
152	The particle si
153	weredetermin
154	technique that
155	distribution of
156	thecolloidaly
157	technique is th
158	preserves the
159	as recovery ca
160	
161	2.5. As and Fe

acterizationand elemental quantification by AF4–ICP-MS zedistribution and elemental composition of the colloid suspensions ed by AF4 (AF2000 model, *Postnova Analytics*) coupled to ICP-MS, a has previously been shown to allow the determination of the size natural colloids generated in contaminated soils and the quantification of mobilized metal(loid)s (Neubauer et al., 2013). An advantage of the AF4 ne mild separation conditions in absence of stationary phase, which original colloid size distribution. The AF4-ICP-MS specifications as well elculations are presented in the Supplementary Material (Tables S2 and S3). eK-edge X-ray absorption spectroscopy (XAS) analyses Arsenic and Fe XAS measurements were done on the colloids (1000-10 nm) isolated from the colloid suspension obtained from the waste-pile and the sedimentationpondafter 20 and 50 minutes of simulated rainfall. Arsenic and Fe K-edge EXAFS spectra were recorded at the bending magnet BM25A beamline at the ESRF-European Synchrotron Radiation Facility (Grenoble, France). Additional As EXAFS spectra from the SP colloids were recorded at beamline 22 of the ALBA Synchrotron Facility (Barcelona, Spain). Methods and reference materials are described in the Supplementary Material(Table S4).

170

162

163

164

165

166

167

168

169

171 3. RESULTS

172 3.1. Characterization of colloid-suspensions

The volumes, pH and electrical conductivity of runoff suspensions collected during the 173 simulated 60-minutes rain events with 28 mm of rainfall are listed in Table 2. Given the 174

collected volumes and the total amount of distilled water applied per rain event, only 175 6% (WP) to 19% (RB) of the applied rainfall contributed to surface runoff, whereas the 176 major fraction (94-81%) infiltrated into the subsurface. The acidity and electrical 177 178 conductivity (EC) of the colloid-suspensions remained almost constant with time. The colloid concentration in the suspensions is summarized in Table 3 and Figure 179 S5(Supplementary Material). In the waste-pile and river-bed, with clay contents of 47% 180 and 26% (Table 1), respectively, the colloid concentration in the suspensions remained 181 182 almost constant with time over the rainfall simulation, while in the sedimentation-pond, with the lowestclay content (12%), the colloid concentration in the colloid-183 suspensiondecreased markedly during the initial 20 minutes of rainfall, and varied 184 around a mean value of 0.26 mg L⁻¹ over the rest of the irrigation period(Figure S5). 185 The fractions of total As in runoff in colloidal form (97%, 0.8% 22% for waste-pile, 186 187 river-bed and sedimentation-pond, respectively) indicated near exclusive As 188 mobilization in colloidal form from the waste-pile, an important colloidal As 189 mobilization from the sedimentation-pond and only dissolved As mobilization from the 190 river-bed (Table 3). The fraction of dissolved Fe was negligible in all cases, indicating near-exclusive association of Fe with suspended colloids. 191 The molar As/Fe ratio in colloids inmobilized form the waste-pile zone was near unity, 192 193 while this parameter was 0.36-0.58 and 0.05 in colloids from river-bed and 194 sediementation-pond, respectively. This disparity pointed to differences in he As mode of association to colloids that warrant examination by XAS. 195 196 Both As and Fe concentrations in the colloid-suspension expressed per mass of colloids decreased at the end of the experiments (50 minutes), although the colloid 197 198 concentrationitself remained constant. Notably, the As concentration in the colloidal phase was about 1.9 (river-bed) to 2.9 (waste-pile and sedimentation-pond) times higher 199

than in the corresponding bulk materials (Tables 1 and 3), indicating that As was enriched in the colloid-size fraction of the waste and sediment materials. Conversely, the Feconcentration in colloids mobilized from the waste-pile wassimilar to the Fe concentration in the waste material, whereas the Fe concentration in colloids mobilized from the river-bed and the sedimentation-pond decreased 5-10 times. This could be partially explained by the formation of Fe-(hydr)oxide coatings (Gomez-Gonzalez et al., 2014) on the surface of aluminum-silicate particles (Table 1) larger than the size of colloids (< 1000 nm) in both the river-bed and the sedimentation-pond.

3.2. Size characterization of the colloids from the sedimentation-pond

The colloid suspensions from the waste-pile and the river-bed could not be measured by AF4. In the waste-pile this was due to the high electrical conductivity of its runoff suspensions (Table 2) which led to irreversible colloids absorption onto the AF4 channel membrane. The colloid concentrations in the runoff suspensions from the riverbed were too low to be determined by AF4 (Table 3, Figure S6). Conversely, the particle size distribution and associated As, Fe and Al concentrations in the colloid-suspensions(20, 35 and 50 minutes) of the sedimentation-pond zone were obtained by AF4-ICP-MS (Figures 1 and S7). Size distribution maxima were 147 nm, 160 nm and 195 nm for the colloid-suspensions obtained at 20, 35 and 50 minutes of rainfall, respectively. The maxima corresponding to the secondary shoulder visible at 20 and 35 minutes and to the defined secondary peak at 50 minutes of rainfall varied little (~460 nm) as estimated by fitting the fractograms to a double-peak lorentzian function (Figure 1a).

maxima, white circle; confidence interval (colour line) at 95% probability). Arsenic and

Fe concentration maxima appeared at 143, 115 and 136 nm for the colloid-suspension obtained at 20, 35 and 50 minutes of irrigation time, indicating a similar behavior of these elements in the colloidal mobilization. The Al concentration maxima appeared at 156, 121 and 166 nm for the same irrigation times.

229

230

231

232

233

234

235

236

237

238

239

240

241

242

243

244

245

246

247

248

225

226

227

228

3.3. Speciation of colloidal As and Fe by K-edge XAS

The As and Fe K-egde XANES spectra of colloids isolated from runoff from the wastepile and sedimentation-pond (Figure S8, Supplementary Material), indicatedAs(V) and Fe(III) as the dominant oxidation states. The As and Fe K-edge EXAFS spectra of colloids from the waste-pile runoff and XANES spectra of colloids from the sedimentation-pond together with the reconstructed linear combination fitting (LCF) spectra are shown in Figure 2and the LCF results are listed in Table 4. In the case of colloids from the sedimentation-pond, only the XANES spectra were evaluated ue to the low quality of the corresponding EXAFS spectra. Arsenic K-edge EXAFS spectra from the waste-pile runoff indicated a major contribution (91–94%) fromscoroditeand minor contribution of As(V) sorbed to ferrihydrite (6–9%) at both irrigation times. Iron K-edge EXAFS spectra also showed a major contribution from scorodite (91–96%) and minor amounts of nontronite (9%) and hematite (4%). Results from shell-by-shell fit analysis of As and Fe K-edge EXAFS spectra of colloids from the waste-pile were in line with a dominant fraction of As and Fe in scorodite (Table S5, details on shell-fit procedure in Supplementary Material). The LCF analysis of the As and Fe K-edge XANES spectra of colloids from the sedimentation-pond on the other hand, revealed a dominant fraction of As(V) sorbed to ferrihydrite and Fe(III) in smectite (50%), nontronite (21–27%) and schwertmannite

(23–29%) (Table 4, Figure 2) suggesting that Fe was mainly contained in clay minerals, and to a lesser extent poorly crystalline Fe(III)-(hydr)oxides.

251

252

253

254

255

256

257

258

259

260

261

262

263

264

265

266

267

268

269

270

271

272

273

249

250

4. DISCUSSION

4.1. Colloidal scorodite as potential nanovector for As dispersion with surface runoff Although scorodite may be thermodynamically unstable when exposed to water containing low Fe and As concentrations, the dissolution of synthetic scorodite over the pH range 5-9 has been shown to be slow(Bluteau and Demopoulos, 2007),andPaktunc and Bruggeman(2010)reported that As solubility in equilibrium with scorodite is lowest at pH 3-4. Our results show that colloidal scorodite mobilized with surface runoffis acting as a mobileAs carrier. Molecular-scale colloidal As speciation confirms the presence of scorodite in the colloids isolated from the waste-pile colloid-suspensions (Tables 4 and S5, Figure 2). These results are in agreement with the molar As/Fe ratio in the colloid-suspension (Table 3) and the XRD analyses (Table 1). The colloidal stability and mobility of scorodite is expected to depend on its surface charge, but respective data is scarce. According to Robins (1987), scorodite formed at pH < 1.8 exhibit a negative surface charge, whereas the formation of a surface Fe(III)-(hydr)oxide may lead to a positive surface charge. Other surface transformation on the colloidal scorodite may also take place affecting the stability of weathered scorodite colloids. From the waste-pile material, As was mainly mobilized in colloidal form, with colloidal As concentrations between 3512 and 4488 µg L⁻¹. Based on the average As concentration in the colloid-suspensions from the waste-pile(Table 3) and the volume of surface runoff, about 6.7 mg of colloidal As were released with the surface runoff from the irrigated area (0.24 m²) during the 60-minutes rainfall simulation, corresponding to

~28mg colloidal As per m²waste-pile surface. Considering that the dumped processing wastes cover about 36 m², roughly 1 g of colloidal As could be released by surface runoff during a natural rainfall event with a duration and intensity similar to the simulation (60 minutes, 28 mm of rainfall). During the simulated rainfall event, 6% of the water volume applied to the waste pile surface was collected as surface runoff (Table 2), suggesting that most water infiltrated into the waste material. Colloidal scorodite could also contribute to vertical transport of As into deeper soil horizons. Indeed, a recent study at the same site (Gomez-Gonzalez et al., 2014)indicated the presence of scorodite at greater depths, although the fraction of scorodite decreased with depth at the expense of As(V) sorbedto Fe(III)-(hydr)oxides. Rainfall intensity and duration affect the runoff/infiltration ratio and this should be considered site-specific. In the studied area, rainfall events with > 10 mm occur on average on 18.5 days per year, and storm events with > 30 mm of precipitation on 2 days per year(Guíaresumida del climaenEspaña (1981-2010), www.aemet.es). This indicates that the potential for continuing As dispersion via colloidal As release and surface runoff is substantial unless remediation actions are taken.

290

291

292

293

294

295

296

297

298

274

275

276

277

278

279

280

281

282

283

284

285

286

287

288

289

4.2. Potential migration of colloidal As into the downstream soils

Arsenic-bearing colloids in runoff suspensions from the river-bed and the sedimentation-pond greatly differ from those identified in the waste-pile zone. Two processes are depicted from our results: First, virtually no colloidal scorodite mobilized fromthe waste-pile appears at the river-bed. At this second experimental zone, colloidal As concentrations are negligible compared to dissolved ones (Table 3). Along the riverbed, a large fraction of As released after scorodite dissolution might be leached through the river-bed itself, and through preferential flow paths in the subsoil, as suggested by

Helmhart et al. (2012). Second, in the sedimentation-pond, where a major part of runoff waters isretained and may remain stagnant over several days before drying out, colloidal As is mostly As(V) associated with ferrihydrite or other Fe(III)-(hydr)oxides. The increasing pH in the river-bed and sedimendation-pond, i.e., with increasing distance from the waste pile may favor scorodite dissolution and As(V) resorption to (amorphous) Fe(III)-(hydr)oxide colloids(Serrano et al., 2015). Ferrihydrite has a high capacity for As(V) uptake(Fritzsche et al., 2011), and their association has been extensively reported(Johnson and Hallberg, 2005). The present study indicates that a fraction of As-bearing Fe(III)-(hydr)oxide colloids contained in the sedimentation-pond as spectroscopically described in the dispersible colloidal fraction by Serrano et al. (2015) are potentially mobilizable and may behave as colloidalAs carriers within runoff generated by strong rain events. The differences in colloid size-distributions derived from AF4-UV-vis (Figure 1a) and AF4-ICP-MS fractograms (Figure 1b) as well as the slight discrepancy in temporal trends of Fe and Al (Figures 1b and S7) may be due to Al-containing clay minerals in the colloidal fraction of the sedimentation-pond runoff (Table 4), whoselaminar structure affects the hydrodynamic diameter of the particles measured by AF4. The use of distilled water as rainfall instead of artificial rainwater with background electrolyte concentrations may introduce some disparity with natural rain events. However, Vazquez et al. (2003) reported ionic rainfall compositions in NW Spain with a range of electrical conductivities similar to the ones obtained in our river-bed and sedimentation-pond runoff suspensions. According to that, we would not expect drastic differences in the use of distilled water rather than a slight enhance of colloids release.

322

323

299

300

301

302

303

304

305

306

307

308

309

310

311

312

313

314

315

316

317

318

319

320

321

5. CONCLUSIONS

The results from this study suggest that colloidal scorodite may be released from mining wastes and act as an As carrier in surface runoff. Similarly, As(V) immobilize by sorption to Fe(III)-(hydr)oxidesin the sediments from the sedimentation-pond may be susceptible to mobilization in colloidal form. Deposited mine wastes from legacy mining activities in Spain, should therefore be isolated to reduce weathering and colloid transport with surface runoff which may result in the dispersion of As in both colloidal and dissolved As form. This should be taken into account when As-rich metallurgical wastes are deposited above ground, because both scorodite and ferryhidrite should not be considered as effective trapping systems of As(V). More research is needed, however, to better quantify the importance of As mobilization with surface and subsurface runoff in colloidal and dissolved form in order to effectively prevent As release to the environment. Also, the application of any capping system aiming to prevent the wastes from erosion could limit the colloids and contaminants migration in the environment. Rainfall simulation experiments and advanced separation techniques such as AF4-ICP-MS are suitable analytical procedures to characterize suspended colloids in runoff and study their potential for colloid-facilitated contaminant release and transport.XAS provides complementary molecular-level insight into the speciation of inorganic contaminants in the colloidal phase.

343

344

345

346

347

348

324

325

326

327

328

329

330

331

332

333

334

335

336

337

338

339

340

341

342

ACKNOWLEDGEMENTS

The Spanish Ministry of Economy and Competitiveness (research project CGL2010-17434) supported this study. M.A. Gomez-Gonzalez was supported by the Ph.D. Spanish FPI fellowship (BES-2011-046461) and by graduate students (EEBB-I-14-08063) programs. XAS measurements on BM25A beamline (ESRF) were supported by

349	the project EV-41; XAS measurements on BL22-CLAESS beamline (ALBA-CELLS)
350	were supported by the project 2012100332. ICP-OES and ICP-MS
351	analyseswereperformed at Servicio de análisis químico (Servicio General de Apoyo a la
352	Investigación-SAI), Universidad de Zaragoza.
353	
354	FIGURE CAPTIONS
355	1– (a) AF4-UV-vis fractograms of colloid-suspension from the sedimentation-pond at
356	20 (black), 35 (red), and 50 minutes (blue). (b) Concentration maxima (white
357	circle)fractogramsand confidence interval (colour line) at 95% probability.Dotted line
358	represents the average size maxima of all elements and times (131 nm of size).
359	2- Arsenic (a) and iron (b) K-edge EXAFS spectra of waste-pile (WP) colloids. Arsenic
360	(c) and iron (d) XANES spectra of sedimentation-pond (SP) colloids.Black
361	lines, experimental data; red lines, LCF results(Table 4 for LCF values).
362	
363	REFERENCES
364	Bian B, Cheng X-J, Li L. Investigation of urban water quality using simulated rainfall in
365	a medium size city of China. Environmental Monitoring and Assessment 2011;
366	183: 217-229.
367	Bluteau MC, Demopoulos GP. The incongruent dissolution of scorodite - Solubility,
368	kinetics and mechanism. Hydrometallurgy 2007; 87: 163-177.
369	Calvo A, Gisbert J, Palau E, Romero M. Un simulador de lluvia portatil de facil
370	construccion. Eds. M. Sala and F. Gallart. Metodos y Tecnicas para la medicion
371	de procesos geomorfologicos 1988; SEG Monogr.: 6-15, Spain.
372	Cerda A, Ibanez S, Calvo A. Design and operation of a small and portable rainfall
373	simulator for rugged terrain. Soil Technology 1997; 11: 163-170.

374	Cornell RM, Schwertmann U. The Iron Oxides: Structure, Properties, Reactions,
375	Occurrences and Uses. Wiley-VCH Verlag GmbH & Co. KGaA, 2004.
376	Cheng H, Hu Y, Luo J, Xu B, Zhao J. Geochemical processes controlling fate and
377	transport of arsenic in acid mine drainage (AMD) and natural systems. Journal
378	of Hazardous Materials 2009; 165: 13-26.
379	Fernandez-Galvez J, Barahona E, Mingorance MD. Measurement of infiltration in small
380	field plots by a portable rainfall simulator: Application to trace-element
381	mobility. Water Air and Soil Pollution 2008; 191: 257-264.
382	Fritzsche A, Rennert T, Totsche KU. Arsenic strongly associates with ferrihydrite
383	colloids formed in a soil effluent. Environmental Pollution 2011; 159: 1398-
384	1405.
385	Garcia-Sanchez A, Alvarez-Ayuso E. Arsenic in soils and waters and its relation to
386	geology and mining activities (Salamanca Province, Spain). Journal of
387	Geochemical Exploration 2003; 80: 69-79.
388	Gomez-Gonzalez MA, Serrano S, Laborda F, Garrido F. Spread and partitioning of
389	arsenic in soils from a mine waste site in Madrid province (Spain). Science of
390	the Total Environment 2014; 500–501: 23-33.
391	Grosbois C, Courtin-Nomade A, Robin E, Bril H, Tamura N, Schäfer J, et al. Fate of
392	arsenic-bearing phases during the suspended transport in a gold mining district
393	(Isle river Basin, France). Science of the Total Environment 2011; 409: 4986-
394	4999.
395	Helmhart M, O'Day PA, Garcia-Guinea J, Serrano S, Garrido F. Arsenic, Copper, and
396	Zinc Leaching through Preferential Flow in Mining-Impacted Soils. Soil Science
397	Society of America Journal 2012; 76: 449-462.

398	Johnson DB, Hallberg KB. Acid mine drainage remediation options: a review. Science
399	of the Total Environment 2005; 338: 3-14.
400	Kretzschmar R, Sticher H. Transport of Humic-Coated Iron Oxide Colloids in a Sandy
401	Soil: Influence of Ca2+ and Trace Metals. Environmental Science &
402	Technology 1997; 31: 3497-3504.
403	Laborda F, Ruiz-Begueria S, Bolea E, Castillo JR. Study of the size-based
404	environmental availability of metals associated to natural organic matter by
405	stable isotope exchange and quadrupole inductively coupled plasma mass
406	spectrometry coupled to asymmetrical flow field flow fractionation. Journal of
407	Chromatography A 2011; 1218: 4199-4205.
408	Morin G, Calas G. Arsenic in soils, mine tailings, and former industrial sites. Elements
409	2006; 2: 97-101.
410	Neubauer E, von der Kammer F, Knorr KH, Peiffer S, Reichert M, Hofmann T.
411	Colloid-associated export of arsenic in stream water during stormflow events.
412	Chemical Geology 2013; 352: 81-91.
413	O'Day PA, Rivera N, Root R, Carroll SA. X-ray absorption spectroscopic study of Fe
414	reference compounds for the analysis of natural sediments. American
415	Mineralogist 2004; 89: 572-585.
416	Paktunc D, Bruggeman K. Solubility of nanocrystalline scorodite and amorphous ferric
417	arsenate: Implications for stabilization of arsenic in mine wastes. Applied
418	Geochemistry 2010; 25: 674-683.
419	Ranville JF, Chittleborough DJ, Beckett R. Particle-size and element distributions of
420	soil colloids: Implications for colloid transport. Soil Science Society of America
421	Journal 2005; 69: 1173-1184.

422	Rauch JN. The present understanding of Earth's global anthrobiogeochemical metal
423	cycles. Mineral Economics 2012; 25: 7-15.
424	Regelink IC, Voegelin A, Weng L, Koopmans GF, Comans RNJ. Characterization of
425	Colloidal Fe from Soils Using Field-Flow Fractionation and Fe K-Edge X-ray
426	Absorption Spectroscopy. Environmental Science & Technology 2014; 48:
427	4307-4316.
428	Robins RG. Solubility and Stability of Scorodite, FeAsO $_4.2H_2O$ - Discussion. American
429	Mineralogist 1987; 72: 842-844.
430	Serrano S, Gomez-Gonzalez MA, O'Day PA, Laborda F, Bolea E, Garrido F. Arsenic
431	speciation in the dispersible colloidal fraction of soils from a mine-impacted
432	creek. Journal of Hazardous Materials 2015; 286: 30-40.
433	Vaughan DJ. Arsenic. Elements 2006; 2: 71-75.
434	Vazquez A, Costoya M, Pena RM, Garcia S, Herrero C. A rainwater quality monitoring
435	network: a preliminary study of the composition of rainwater in Galicia (NW
436	Spain). Chemosphere 2003; 51: 375-386.
437	Voegelin A, Weber FA, Kretzschmar R. Distribution and speciation of arsenic around
438	roots in a contaminated riparian floodplain soil: Micro-XRF element mapping
439	and EXAFS spectroscopy. Geochimica Et Cosmochimica Acta 2007; 71: 5804-
440	5820.
441	Wicke D, Cochrane TA, O'Sullivan AD. Atmospheric deposition and storm induced
442	runoff of heavy metals from different impermeable urban surfaces. Journal of
443	Environmental Monitoring 2012; 14: 209-216.

Table Click here to download Table: Rainfall Simulation_Tables_Revised.docx

Table 1 – Physical and chemical properties of bulk samples

	mieur properties or t	Samples	
Characteristics	WP	RB	SP
pH ^a	3.3	4.3	6.1
$E.C. (\mu S cm^{-1})^b$	2320	93.7	110
ECEČ (eq kg ⁻¹) ^c TOC (%) ^d	0.79	0.11	0.05
TOC (%) d	0.03	0.24	0.20
Soil texture (%) ^e			
Sand	53	30	70
Silt	-	44	18
Clay	47	26	12
Element concentration (mg k	$(g^{-1})^f$		
Al	1100	422	257
Fe	25230	45530	16980
Cu	1390	99	52
Zn	151	82	73
As	14390	620	94
Pb	465	29	22
Mineralogical composition ^g			
Major phases	Scorodite	Illite	Illite
-	Gypsum	Kaolinite	Microcline
		Quartz	Kaolinite
			Albite
Minor phases	Quartz	Muscovite	Quartz
	Hematite	Albite	Montmorillonite

^aPotassium chloride (KCl, 3 mol L^{-1}) was used as electrolyte solution for pH measurements

Table 2 – Physical characteristics of the collected runoff suspensions^a

Sampling location	Runoff time	Runoff volume ^b	рН	Electrical Conductivity
location	min	mm	-	$\mu S cm^{-1}$
WP	0-10	0.27	3.90	1567
	10-20	0.23	4.08	1532
	20-30	0.27	4.16	1580
	<i>30-40</i>	0.24	4.04	1640
	40-50	0.30	4.14	1585
	50-60	0.32	4.20	1548
	Total surface runoff ^c	1.63 (6%)		
RB	0-5	0.45	4.54	198
	5-10	0.45	4.60	164
	10-15	0.47	4.63	143
	15-20	0.45	4.64	133
	20-25	0.44	4.60	147
	25-30	0.43	4.52	163
	30-35	0.37	4.57	169
	35-40	0.40	4.63	160

 $^{{}^{}b}E.C. = Electrical conductivity$

 $^{^{}c}ECEC = Effective\ Cation\ exchange\ capacity\ as\ the\ sum\ of\ Ca,\ Mg,\ Na,\ K\ and\ Al\ (Shuman\ 1990)$ $^{d}TOC = Total\ organic\ carbon\ determined\ by\ wet\ digestion\ (Walkley\ and\ Black\ 1934)$

^eTexture was determine by the pipette method after removing soil organic matter (Gee and Bauder 1982)

fPseudo-total concentration measured by ICP-OES after agua regia + microwave digestion (Chen and Ma 2001)

^gDetermined by qualitative XRD analysis. Major phases are those with more than 15 % contribution to the total composition

	40-45	0.46	4.56	159
	45-50	0.44	4.54	156
	50-55	0.48	4.68	136
	55-60	0.49	4.54	157
	Total surface runoff c	5.33 (19 %)		
SP	0-5	0.28	6.24	53
	5-10	0.31	6.26	52
	10-15	0.30	6.44	44
	15-20	0.37	6.33	43
	20-25	0.39	6.41	40
	25-30	0.42	6.18	41
	30-35	0.44	5.96	46
	35-40	0.46	5.90	38
	40-45	0.48	5.94	37
	45-50	0.45	6.25	33
	50-55	0.44	6.34	30
	55-60	0.48	6.43	23
	Total surface runoff c	4.81 (17 %)		
// D : .11 1	1 . 1 . 0.11			

Table 3 –As and Fe concentrations of colloid (≤ 1000 nm) suspensions (CS) and dissolved fractions(DF) ≤ 10 nm (\pm standard deviation, n=3 a)

Sample	Time	Fraction _	As	Fe	Colloid concent.d	Molar As/Fe	As concent.	Fe concent.
Sample	Time	Traction —	µg I	1 b,c	mgL ⁻¹	ratio	mg kg c	olloid ^{-1 e}
WP	20 min	CS	4500±5	3230±303				
		DF	128±7	< QL	109	1.04	$41 \cdot 10^{3}$	$30 \cdot 10^3$
		Colloids	4372 ^f	3230				
WP	40 min	CS	4600±23	3480 ± 215				
		DF	112 ± 2	< QL	95	0.98	$48 \cdot 10^3$	$36 \cdot 10^3$
		Colloids	4488	3480				
WP	50 min	CS	3590±18	2660±382				
		DF	78±2	< QL	102	1.01	$35 \cdot 10^3$	$26 \cdot 10^3$
		Colloids	3512	2660				
RB	20 min	CS	281±6	588±303				
		DF	298±4	< QL	200	0.36	1410	2940
		Colloids	-	588				
RB	35 min	CS	262±6	587±185				
		DF	260±6	<ql< td=""><td>207</td><td>0.33</td><td>1270</td><td>2840</td></ql<>	207	0.33	1270	2840
		Colloids	2	587				
RB	50 min	CS	146±7	186 ± 206				
		DF	148 ± 8	< QL	154	0.58	950	1210
		Colloids	-	186				
SP	20 min	CS	131±5	1870±253				
		DF	100±4	< QL	399	0.05	330	4690
		Colloids	31	1870				
SP	35 min	CS	128±6	1990±290				
		DF	106±5	< QL	437	0.05	290	4550
		Colloids	22	1990				
SP	50 min	CS	96±6	1060 ± 348				
		DF	72±6	< QL	427	0.05	200	2200
		Colloids	24	1060				

^aStandard deviation was calculated by triplicate determination of the same aliquot

^aDistilled water was used in the rainfall simulation tests
^bVolume of runoff collected during the rainfall simulation test at waste-pile (WP), river-bed (RB) and sedimentation-pond (SP)

^cDefined as the sum of all volumes collected during the experiment. The percentage of water collected respect to the total water pumped during the experiment (28 mm) is presented in parenthesis

Table 4 – Linear combination fit results for As and FeK-edge EXAFS or XANES spectra

Sample	Scor	odite ^b	As sor Ferrih		R - Factor ^e	$red \chi^{2f}$
-	% ^c	ΔeV^d	%	∆eV	- Factor	
WP 20 min ^g	94.2	-	5.7	-	0.0249	0.8562
WP 50 min ^g	91.3	-	8.7	-	0.0199	0.6527
SP 20 min ^h			99.7	- 0.3	0.0116	0.0062
SP 50 min ^h			99.6	- 0.4	0.0125	0.0068

Fe EXAFS/XANES a

Sample	Scorodite ^b		Scorodite ^b Nontronite ^b Hematite ^b		Smectite ^b		Schv man	wert- mite ^b	R factor	red χ ²		
	%	∆eV	%	∆eV	%	∆eV	%	∆eV	%	∆eV	lactor	
WP 20 min ⁱ		-	8.6	-							0.0517	0.665
WP 50 min ⁱ	95.7	-			4.2	-					0.0505	0.619
SP 20 min j			27.1	0.1			50.3	- 0.1	23.3	0.2	0.0003	0.0001
SP 50 min ^j			20.9	0.3			50.4	- 0.1	29.3	- 0.1	0.0003	0.0001

^aEXAFS spectra of waste-pile(WP) colloids and XANES spectra of sedimentation-pond (SP) colloids were analyzed by LCF analyses

^bElementalconcentration was expressed on µg of metal per liter of colloid suspension (CS)

^cQuantification limits (QL) for ICP-MS measurement of As was 0.1 μ g L^{-1} and 73 μ g L^{-1} for ICP-OES measurements of Fe

^dMass of colloids calculated according to Plathe et al. 2010

^eMilligrams of As and Fe per kilogram of colloids (1000-10 nm)

 $[^]f$ As and Fe concentrations of colloids (1000-10 nm) were defined as the difference between the concentration of the CS and the DF

^bThe scorodite reference spectrum was obtained from Savage et al. (2005). The As(V) sorbed to ferrihydrite spectrum was obtained from Root et al. (2009). The nontronite (NAu-1) spectrum was obtained from Gorski et al. (2013). The hematite and smectite spectra were obtained from O'Day et al. (2004). The schwertmannite spectrum was measured by S. Hayes in the BM 4-1(SSRL, Stanford, USA, unpublished) (Table S4, Supplementary Material)

^cFittings were not constrained to sum 100 %

^dFittingvariations of XANES LCF given by the software Athena (Ravel and Newville 2005)

^eNormalized sum of the squared residuals of the fit $[R = \sum (data-fit)^2 / \sum data^2]$

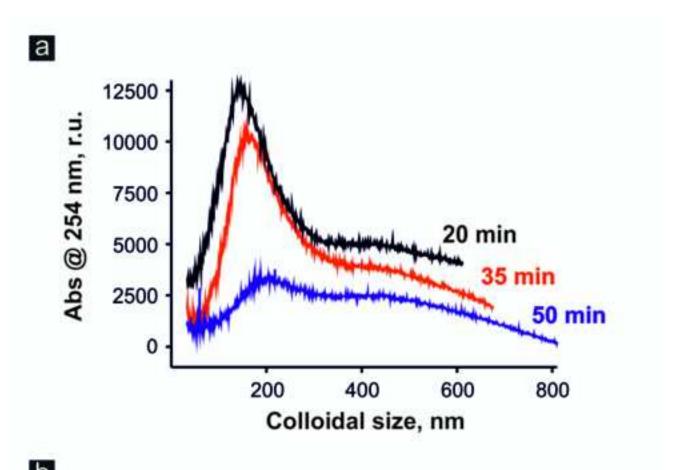
^fGoodness-of-fit was assessed by the χ^2 statistic [=(Ffactor)/(no. of points - no. of variables)]

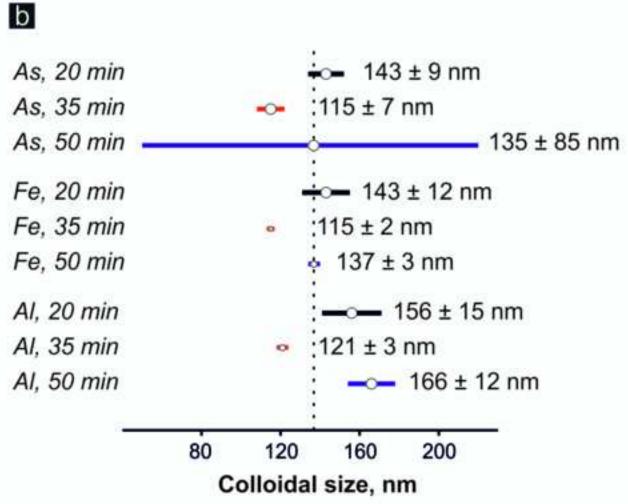
^gLCF EXAFS range: 2 – 11 Å, spectra measured at ESRF (Grenoble, France)

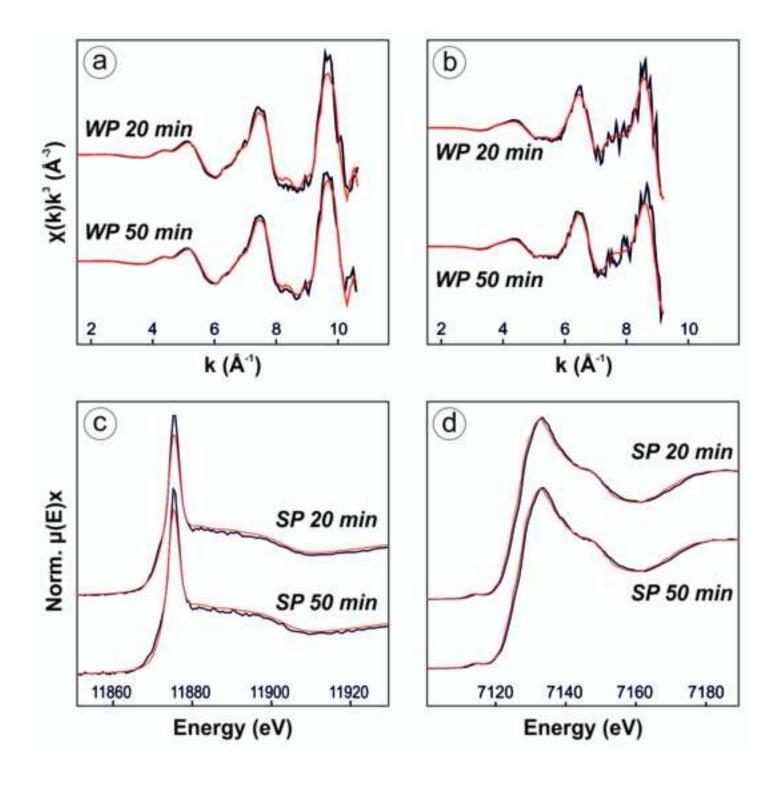
^hLCF XANES range: 11855 – 11935 eV, spectra measured at ALBA (Barcelona, Spain)

ⁱLCF EXAFS range: 2 – 9 Å, spectra measured at ESRF (Grenoble, France)

^jLCF XANES range: 7105 – 7185 eV, spectra measured at ESRF (Grenoble, France)







Supplementary Material
Click here to download Supplementary Material: Rainfall Simulation_SM_Revised.pdf

