

Investigation of small-scale processes in the rhizosphere of *Lupinus albus* using micro push-pull tests

K. Knecht · B. Nowack · M. H. Schroth ·
M. J.-F. Suter · R. Schulín

Received: 20 June 2013 / Accepted: 9 January 2014 / Published online: 31 January 2014
© Springer International Publishing Switzerland 2014

Abstract

Background and Aims Rhizosphere processes affect the mobility, phytoavailability and toxicity of solutes in soil. To study reactions in the rhizosphere under quasi in situ conditions, we recently developed the “micro push-pull test” (μ PPT) method, combining micro-suction cups with the principle of the “push-pull test” method known from groundwater applications. Here we report the

application of μ PPT to investigate rhizosphere reactions in situ, i.e. degradation of deuterated citrate (citrate-d4) in the rhizosphere of *Lupinus albus* grown in sand-filled rhizoboxes.

Methods In a μ PPT, a solution containing reactive (citrate-d4) and non-reactive solutes (bromide) is injected into a porous medium and shortly thereafter, the pore water solution is re-extracted from the same location. Concentration (“breakthrough”) curves of extracted reactants can be compared to those of the non-reactive solute, allowing the determination of reaction rates. We applied the μ PPT in rhizoboxes with *Lupinus albus* and sampled different types of micro-habitats: bulk soil, rhizosphere of normal roots and rhizosphere of cluster roots of different ages.

Results Breakthrough curves of citrate-d4 varied considerably between tests adjacent to cluster roots and normal roots, and in bulk soil. Degradation of citrate-d4 in bulk soil and adjacent to normal roots was below detection, while we found strong degradation of citrate-d4 adjacent to 4 to 5-days old cluster roots. In situ degradation rate constants for citrate-d4 around cluster roots were found to be in the range from 0.38 to 0.71 h⁻¹.

Conclusions We successfully applied the μ PPT to the rhizosphere. The μ PPT is useful to investigate local processes in microcosms and to monitor processes also over time (e.g., during cluster-root development) due to its non-destructive nature.

Keywords Citrate · Microbial degradation · Micro push-pull test · Modeling · Rhizosphere · Suction cup

Responsible Editor: Timothy George.

K. Knecht · B. Nowack · R. Schulín
Institute of Terrestrial Ecosystems, ETH Zurich,
8092 Zurich, Switzerland

K. Knecht
e-mail: kajsa.knecht@bluewin.ch

R. Schulín
e-mail: schulín@env.ethz.ch

B. Nowack (✉)
Swiss Federal Laboratories for Materials Science and
Technology (EMPA),
Lerchenfeldstrasse 5, 9014 St. Gallen, Switzerland
e-mail: nowack@empa.ch

M. H. Schroth
Institute of Biogeochemistry and Pollutant Dynamics,
ETH Zurich,
Zurich, Switzerland
e-mail: schroth@env.ethz.ch

M. J.-F. Suter
Eawag: Swiss Federal Institute of Aquatic Science and
Technology,
Dübendorf, Switzerland
e-mail: Marc.Suter@eawag.ch

Introduction

The rhizosphere, the zone of soil under direct influence of active plant roots, differs in many aspects from bulk soil. For example, root growth and water uptake by plants directly alter physical properties of the rhizosphere. Therefore, flow and transport processes can vary in the rhizosphere and strongly differ from those in bulk soil. Furthermore, root activities such as nutrient absorption, respiration, and exudation strongly influence chemical properties of the rhizosphere (e.g. concentration of nutrients, of toxic elements, of complexing compounds, pH, pO_2 and pCO_2). Moreover, root exudation stimulates microbial activity, resulting in strong gradients in chemical potentials and solute concentrations (e.g. Hinsinger et al. 2009; Vetterlein and Jahn 2004). Thus, the rhizosphere is a unique micro-habitat and most of all a zone of particularly intense turnover and transformation processes. To understand the ecological role and functioning of the rhizosphere, it is highly important to understand rhizosphere processes and the factors governing them.

Much progress has been made during the past decades in assessing the spatial variability in soils and especially in the rhizosphere (Gregory and Hinsinger 1999; Luster et al. 2009). In the 1990s, small-scale soil-solution investigations started with the miniaturization of sampling probes, which enabled non-destructive in situ observation of soil-solution chemistry (Göttlein et al. 1996; Vetterlein et al. 1993). Later, micro-suction cups were used successfully in conjunction with rhizoboxes that allow observation of root-system development through a transparent front plate and thus permitted the targeted sampling of soil solution adjacent to roots (Dessureault-Rompré et al. 2006, 2007; Dieffenbach et al. 1997; Göttlein et al. 1999; Puschenreiter et al. 2005; Schreiber et al. 2011). This allowed the determination of solute concentrations in the rhizosphere at high spatial and temporal resolution and provided information about small-scale heterogeneities with respect to solute concentrations. Nevertheless, the underlying processes are often not yet adequately understood: So far, micro-scale processes were mostly studied under rather artificial conditions (e.g., pot and column experiments, hydroponics) (Luster et al. 2009) and little is known about the mobility and degradation of solutes (e.g. exudates) under real rhizosphere conditions. New non-destructive methods to investigate rhizosphere processes in situ are therefore needed. One example is the use of optodes for non-destructive

analysis of pH and oxygen in the rhizosphere (Blossfeld et al. 2011).

On a larger scale, single-well injection-withdrawal tests, called push-pull tests, have been used extensively for the investigation of chemical, physical and biological phenomena in aquifers (Drever and McKee 1980; Istok et al. 1997). In a push-pull test, a solution containing non-reactive and reactive tracers is injected (pushed) into an aquifer. The test-solution/groundwater mixture is then extracted (pulled) from the same location. The concentration (“breakthrough”) curves of the extracted reactants are compared to those of the tracers, which allows to determine the rates at which the reactants are eliminated from the solution.

The idea underlying the study presented here is the miniaturization of the push-pull technique by using micro-suction cups to investigate small-scale processes in the rhizosphere. The new technique is expected to yield not only concentrations of solutes, but also reaction rates for processes such as the biodegradation of exudates or the ligand-controlled dissolution of minerals. In a previous study, we showed that the push-pull technique can be miniaturized using micro-suction cups (Knecht et al. 2011). By injecting about 250 μl of a solution containing different non-reactive and reactive tracers into a sand-filled, water-saturated thin-slab chamber and then extracting about 850 μl of the pore water from the same point with a slow and constant flow rate, we obtained breakthrough curves that were well predicted by simulations based upon the advection-dispersion equation (White and Oostrom 2000). In a similar experimental system, we subsequently adopted the μPPT method to unsaturated conditions and showed that it worked to a soil water saturation as low as 34 % (Knecht et al. 2013).

All these previous tests were performed in the absence of plant roots. Root activity changes the relative hydraulic conductivity, which is the main limiting factor for the μPPT method under unsaturated conditions, not only by altering the water content but also because of compaction and reorientation of soil particles around them. In addition, microbial activity can have a strong influence on flow conditions (during a μPPT experiment). Microorganisms colonizing the pore space can clog pores by proliferation of cells, production of extracellular polymers, release of gaseous byproducts or microbially mediated accumulation of insoluble precipitates (“bioclogging”) (Baveye et al. 1998). Soil hydraulics may not only be affected by changes in porosity, but also by changes in the viscosity of the soil solution. All

these changes can be local and result in a non-uniform porous medium characterized by different local porosity domains with highly contrasting permeability. Under unsaturated conditions, these heterogeneities can cause hysteresis effects during a μ PPT experiment, thus the sampled pores during extraction may not be entirely identical with those affected during injection of the test solution. Another implication might be caused by preferential flow, which includes “all phenomena where water and solutes move along certain pathways, while bypassing a fraction of the porous matrix” (Simunek et al. 2003).

The aim of this work was to adapt the new μ PPT method to a porous medium colonized by plants and to evaluate its potential usefulness for studying rhizosphere processes. For this purpose we performed μ PPT in the surroundings of roots of *Lupinus albus* grown in sand-filled rhizoboxes. To show the potential of the new method, we studied the biodegradation of citrate as a well-known model process. Citric acid is a low-molecular-weight organic acid, which is exuded by the roots of many plants and thus is a representative substrate for microbial activity in the rhizosphere. *Lupinus albus* was chosen as a model plant because it is well known for exuding particularly high amounts of organic compounds (like citrate) from so-called cluster or proteoid roots (Dinkelaker et al. 1995).

To investigate degradation, we used deuterated citrate (citrate-d4), which was injected both into bulk soil and adjacent to roots. As a non-reactive tracer bromide was injected simultaneously to obtain information about the hydraulic properties of the pore space into which the test solution was injected. The extracted transient solute concentration curves were numerically modeled. To account for changes in the tortuosity and permeability at sampling locations, hydraulic parameters were fitted when necessary. In addition, potential in situ degradation rates for citrate-d4 around cluster roots were determined. Background concentrations of citrate (exuded by the plants) were monitored to gain information about the exudation activity of the roots.

Materials and methods

Rhizoboxes

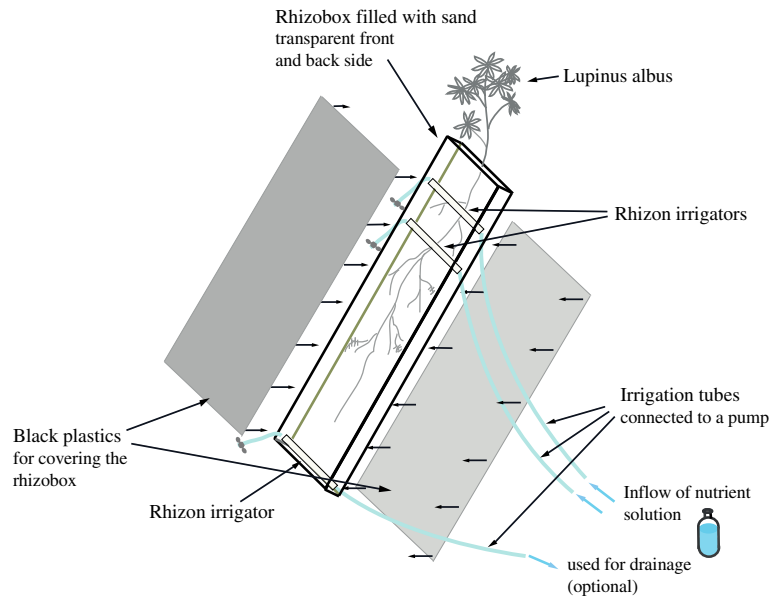
The rhizoboxes used in this study were adapted from the type used by Dessureault-Rompré et al. (2006). The

boxes had a length of 60 cm, a width of 15 cm and a thickness of 1 cm (inner volume: 900 cm³; Fig. 1). Front and back sides were made of acrylic glass plates to allow the observation of root growth and to permit the targeted insertion of micro-suction cups. The two transparent plates were fixed into a frame made from polyvinylchloride (PVC) and sealed with silicon to prevent water leaking from the box. Three wicks made of a hollow polyethersulfone fiber (Rhizon irrigators, Rhizosphere research products, Wageningen, Netherlands) were installed, one each at 5 and 10 cm depth, enabling subsequent irrigation of the system, and one at 58 cm depth, allowing for drainage at the bottom. Each rhizobox was filled with quartz sand (40/50 Accusand, Schroth et al. (1996)). The mean particle diameter was 0.359 mm and the sphericity 0.9, the particle size distribution is shown in Schroth et al. (1996), > 99 wt.% of particles are in the range 0.25–0.5 mm. The sand contained trace concentrations of metals and had an organic matter content of 0.3 g kg⁻¹. For further chemical and physical characteristics we refer to Schroth et al. (1996). To pack the sand as homogeneously as possible, we used a wide-neck falling funnel containing three sieves to randomize particle motion during falling. The bulk density of the sand packings was about 1.73 g cm⁻³ in all rhizoboxes.

Plant cultivation

Seeds of *Lupinus albus amiga* (“Weissblühende Lupinie”, Otto Hauenstein Samen, Rafz, Switzerland) were surface-sterilized in a 10 % hydrogen peroxide solution (Liang and Li 2003) and then germinated in pots filled with sand. After a week, three healthy plants were transplanted into three rhizoboxes described above (one plant per rhizobox) and grown in a climate chamber under controlled conditions on a 16 h (22 °C)/8 h (15 °C) day/night cycle at 60 % humidity. The rhizoboxes were tilted at an angle of 30° to force the roots to grow along the front plate. To prevent light interference on root growth, the rhizoboxes were covered with dark plastic sheets (see Fig. 1). The cover was only removed to monitor root development and for the μ PPT experiments. The rhizoboxes were percolated with a nutrient solution (10 % strength modified Hoagland solution according to Tandy et al. (2006)) by means of a peristaltic pump, keeping the sand in the boxes moist throughout the experiments. After 10 days of growth, the phosphate concentration of the irrigation

Fig. 1 Schematic of a rhizobox with a growing plant. For cultivation, rhizoboxes were tilted at an angle of 30°; rhizoboxes were in an upright position during micro push-pull tests



solution was reduced to 5 μM to induce the formation of cluster roots and to maximize citrate exudation (Johnson et al. 1996; Keerthisinghe et al. 1998; Neumann and Römheld 1999).

Emergence of cluster roots

When cluster roots became distinguishable from normal roots, the cluster rootlets had already reached a length of approximately 2 to 3 mm. It was not possible in this study to clearly distinguish cluster roots already at an earlier stage of their development due to their similarity in color to that of the quartz sand. Furthermore, the roots did not always grow close to the transparent front plate. Based on the investigations of Watt and Evans (1999b), we determined the day of “visible” emergence to be after 2 to 3 days of development. For pictures of cluster roots of *Lupinus albus* we refer to Watt and Evans (1999b).

Estimation of water saturation using light transmission

To obtain spatially resolved information about the water saturation in the medium around the roots and in bulk soil, we used visible light transmission combined with image analysis. We built a system similar to that described by Tidwell and Glass (1994). The rhizoboxes were installed in front of a light box (Prolite 5,000 HiQ, Kaiser Fototechnik, Buchen, Germany). A digital

single-lens reflex camera (Pentax K 10D, Lens: Pentax 18–55, Pentax Europe GmbH, Hamburg, Germany) was fixed at a distance of 50 cm from the chamber to record the intensity of the light transmitted through the porous medium. The light box was covered with black cardboard/plastic with an opening just large enough that only a selected area of 10×10 cm of the rhizobox was illuminated. For analysis, the images (raw image format) were converted to gray scale (Adobe Photoshop CS3 Extended). Gray-scale values were calculated from the pixels included in 3 to 4 rectangles of approximately 1×1 mm size each (ca. 590 pixels in total) selected in the close surrounding of the suction cup, and mean values were calculated. This resolution allowed to detect even small-scale heterogeneities in water saturation around a sampling location. Before taking a picture of a rhizobox with roots, a picture of a dry “control” sand pack was recorded. The pictures of the dry “control” sand pack were used for light intensity offset correction, eliminating possible time dependent heterogeneities in the lighting.

Gray values were calibrated to water saturation using a sand pack in a small thin-slab chamber of 10×10×1 cm (identical acrylic glass and PVC-frame as used for rhizoboxes), as described in detail by Knecht et al. (2013). Pictures were taken at dry and saturated conditions and after varying the capillary pressure head in several steps from 0 to 60 cm H₂O by means of a hanging water column. Again, before each picture, an

additional dry “control” sand pack was recorded. Gray-scale values were averaged along a distinct straight line of approximately 3.5 cm (899 pixels) lateral from the center of the sand pack for each picture. To relate the measured gray-scale values to volumetric water saturation, we used the water retention curve determined for the same sand by Schroth et al. (1996).

Micro push-pull test experiments

Micro push-pull tests were carried out in three different types of “micro-habitats”: bulk soil (>2 cm distance from roots), rhizosphere of normal roots, and rhizosphere of cluster roots (~1 mm distance from normal or cluster root). Cluster roots were selected at different stages of development. Preparations for an experiment started as soon as cluster roots became visible in the lower part of the rhizoboxes. To reduce water flow towards the root during the μ PPT, transpiration of the plants was limited by covering the plants with a transparent plastic bag 1 day before a test was started. Approximately 10 h before the test, irrigation was stopped to limit vertical water flow during the μ PPT. One to three h before the experiments, pictures were taken to determine the water content in the sand packs as described above. Then, holes were drilled through the front plate adjacent to cluster roots, normal roots and in bulk soil by means of a hand drill, and suction cups were installed as described by Knecht et al. (2011). They were made of 1-cm-long ceramic capillaries (pure aluminum oxide, effective length ~0.85 cm; 1.2 mm o.d., 0.6 mm i.d.; PI Ceramic, Lederhose, Germany) and had a total dead volume of <6 μ L. One tip of the capillary was closed with hot glue, while the other one was connected via a small ETFE tube (Omnilab, Mettmenstetten, Switzerland) to a peristaltic pump (BVP, Ismatec, Glattbrugg, Switzerland). To determine pH, a drop of soil solution (about 70 μ L) was collected through additionally drilled holes in bulk soil and in the rhizosphere of the selected roots by inserting a short piece of a small ETFE tube connected to a syringe. To limit disturbance around the suction cups, the samples from the rhizosphere for pH measurement were taken on the side of the roots opposite to the suction cups. The pH of the sampled solution was measured using an ion sensitive field effect transistor electrode (ISFET IQ125, IQ scientific Instruments, California, USA). No samples for pH measurements were taken on the second day of experiments

to avoid interference with the sensitive experimental design after the micro-suction cups had been installed.

A total of 24 μ PPT were performed in rhizoboxes moved to an upright position (no tilt): 4 in bulk soil, 7 adjacent to normal roots and 13 close to cluster roots. The tests were always performed in the afternoon, when peak values for citrate exudation from cluster roots can be expected (Dessureault-Rompré et al. 2007). In one rhizobox, tests were repeated after 24 h using the same suction cups. We defined grey-brown/dark-brown colored cluster roots as senescent cluster roots (cf. Neumann et al. 1999). Within this work the following replicate samples were therefore collected: With one suction cup installed at one location and one μ PPT, we collected a sequence of about 10–15 samples for this one location and time of injection. Replicates refer to i) different sampling locations in different rhizoboxes and different roots and ii) repeating a μ PPT at exactly the same location at a later time.

The test solution contained citrate-d4 (5 mM, citric acid-2,2,4,4-d₄, isotopic purity 98 %, ReseaChem GmbH, Burgdorf, Switzerland) as reactive tracer and bromide (1 mM, potassium bromide p.a., Merck, Darmstadt, Germany) as non-reactive tracer.

The pH of the test solution was adjusted to the measured pH values of the respective sampling location (using 1 M NaOH or 1 M HCl). The injected volumes were about 120 to 160 μ L at rates of 2.4 to 3.0 μ L min⁻¹ (injection time: 0.83 to 0.95 h). In 7 tests the extraction did not yield enough soil solution for further analysis. In the other tests, the extraction volumes were between 210 and 500 μ L after an extraction time of approximately 2.8 to 3.4 h. The extraction rates varied between 2.3 and 2.6 μ L min⁻¹ (for four experiments, the rates were \leq 2.1 μ L min⁻¹). The rest phase between injection and extraction was 0.03 to 0.14 h (time needed to prepare peristaltic pump, tubes and sample vials for extraction phase). Information about the exudation activity of the roots was gained by determination of non-labeled citrate concentrations in the extracted soil solution samples.

Preliminary experiments were performed with an additional rhizobox, using Acid Red 1 (a red dye, Azophloxine for Microscopy, Fluka AG, Switzerland) as tracer (1.47 mM, injection volumes: 200 μ L, injection rates: ~2.5 μ L min⁻¹, injection time: 1.7 h) to visualize the spatial distribution pattern of the injected test solution during a μ PPT (not shown).

Analysis of solution samples

All solution samples were collected in autosampler vials with 200 µl inserts (Chromacol, Herts, United Kingdom). About 20 µl of formaldehyde (>36.5 % in water, Fluka AG, Switzerland) was added to each vial before extraction to prevent microbial degradation (cf. Dessureault-Rompré et al. 2006). Each sample was diluted with ultrapure water (TKA-GenPure, Huber & Co. AG, Reich, Switzerland) in order to get sufficient volume for analysis. Sample vials were hermetically sealed and stored at 4 °C. Analysis was carried out within 25 days. Bromide was analyzed by means of ion chromatography using a DX-320 ion chromatography system equipped with an electrical conductivity detector and an EG40 eluent gradient generator (Dionex IC20, column: AS11-HC, eluent generator EG40: potassium hydroxide (1–60 mM), flow: 1.5 ml/min; Dionex AG, Olten, Switzerland).

Citric acid (sodium citrate monobasic anhydrous, Sigma-Aldrich, Buchs, Switzerland) and citric acid-2,2,4,4-d₄ for internal standard were quantified in the sample extracts using a HP Series 1,100 high-performance liquid chromatograph (Hewlett-Packard, Waldbronn, Germany), equipped with an online vacuum degasser (DG4, Henggeler Analytic Instruments, Riehen, Switzerland), a binary high-pressure gradient pump, an autosampler, and a heated column compartment (column temperature was kept at 25 °C). Separation was achieved with a Macherey Nagel AG, CC 125/2 Nucleosil, 120–3, C18 column. A 5 µL aliquot of sample extract was injected and isocratically eluted off the column with 0.2 mL min^{−1} of 100 % nanopure water for 6 min.

A triple quadrupole mass spectrometer (API4000, Applied Biosystems, Rotkreuz, Switzerland) was operated in negative ion mode electrospray. Ionization efficiency was increased by post-column addition of a 2 % ammonia solution at 30 µL/min using a Bischoff 2,200 micro-HPLC pump (Bischoff GmbH, Leonberg, Germany). Nitrogen gas was used as drying and nebulizer gas. Highest sensitivity was achieved with the spray voltage set to −3,800 V, the source temperature to 390 °C, entrance potential at −15.0 V and the declustering potential to −35.0 V. The transition from m/z 191.6 to 112.3 was monitored with a collision energy of −15 V and m/z 191.6 to 87.0 with a collision energy of −25 V for citric acid, and m/z 195.2 to 113.1 with a collision

energy of −20 V and m/z 195.2 to 89.1 with a collision energy of −25 V for citric acid-d₄.

Modeling

The transport behavior of the tracers was simulated as described by Knecht et al. (2013) using the model STOMP (Subsurface Transport Over Multiple Phases, White and Oostrom (2000)). To account for the potential influence of microorganisms and root growth on flow and transport in the sand pack, the values of the porosity (n) and longitudinal dispersivity (α_L) were adjusted by weighted nonlinear least-squares curve-fitting of the simulated tracer curves to the experimental bromide curves (fitting range for n : 0.30–0.35; fitting range for α_L : 0.4–14 mm). The weights were adjusted according to the analytical reliability of the measured solute concentrations and the information content of the profile (increased weights for the first 1 to 2 h of extraction phase). To account for heterogeneity in the (initial) water content distribution (information gained from pictures of visible light transmission), especially around roots, and for differences in the hydraulic properties between rhizoboxes with plants and those with sterile quartz sand used for water content calibration (e.g. differences in porosity due to root growth and microbial activity in the planted rhizoboxes), also the initial hydraulic pressure (P_i) was fitted within a given range of values (fitting range: about ± 5 mm from estimated mean values). Values for the diffusion coefficients were taken from the literature as $D_{\text{Bromide}} = 2.08 \times 10^{-9} \text{ m}^2 \text{ s}^{-1}$ and $D_{\text{Citrate}^{3-}} = 6.23 \times 10^{-10} \text{ m}^2 \text{ s}^{-1}$ (Buffle et al. 2007). Microbial degradation of citrate was described as a first-order reaction. The reaction rate coefficient was determined by weighted nonlinear least-squares curve-fitting of the model to the measured citrate-d₄ concentrations of the extracted solution samples.

Results

Non-reactive tracer behavior

We performed a total of 24 µPPT in the presence of living plant roots. In all experiments, it was possible to inject a small volume of about 120 to 160 µl of test solution. In 17 experiments it was possible to extract 8–10 samples of about 20 to 80 µl volume each. The water saturation of the sampled soil domains was estimated

using the light transmission method to range between 0.7 and 0.9, with small-scale variations especially around cluster roots. In 7 experiments, it was not possible to extract enough soil solution for the chemical analysis of the tracers. In three of these cases the reason for this was too low hydraulic conductivity due to a water saturation of <0.5 at the sampling locations. In the other four experiments (estimated water saturations ≥ 0.65) hydraulic contact was lost between suction cup and porous medium.

Typical examples of the experimental bromide (breakthrough) curves are shown in Fig. 2. The measured bromide concentrations decreased exponentially from the beginning of the extraction phase. Table 1 gives an overview of the best-fit values of the parameters determined by curve-fitting and the corresponding root mean squared errors (RMSE). For 13 experiments, the $\text{RMSE}_{\text{bromide}}$ was between 0.02 and 0.07. The cumulative mass recovery for these experiments was between 52 and 68 %. For the remaining four μPPT experiments, the $\text{RMSE}_{\text{bromide}}$ was comparably high with values between 0.11 and 0.18 and less than 34 % of the injected bromide could be recovered after about 3 h of extraction. A representative example of breakthrough curves for bromide for each case is shown in Fig. 2: a) $\text{RMSE}_{\text{bromide}}=0.02$; b) $\text{RMSE}_{\text{bromide}}=0.12$. The root mean-squared errors indicate a good model fit in the case of curve (a) while the fit was not satisfactory in case (b). In all four experiments with high $\text{RMSE}_{\text{bromide}}$, the measured concentration of bromide decreased faster than in the best-fit simulation.

Preliminary experiments (data not shown) with Acid Red 1 dye showed strong gravity-driven flow of the

injected test solution during the first 2 h of the tests for rather dry conditions with a water saturation <0.45 . In tests with higher water saturation (water saturation >0.7), the front of the injected test solution showed in most cases more or less circular shapes around the suction cups, but in some cases, the test solution showed uneven flow patterns during injection (e.g. distinct rapid flow in only some parts around the suction cup or along a root).

Reactive tracer behavior

After adapting the hydraulics of the respective habitats in the numerical simulations to fit the non-reactive tracer data, theoretical extraction breakthrough curves were calculated with the assumption of no microbial degradation for citrate-d4 with the diffusion coefficient for citrate and with the identical hydraulic properties for flow and transport conditions (for experiments with $\text{RMSE}_{\text{bromide}} \leq 0.07$). Figure 3 shows results from tests in bulk soil, adjacent to normal roots and adjacent to cluster roots for two different rhizoboxes each. In most cases, the simulated curves described the measured citrate-d4 concentrations reasonably well in the early phases of extraction, but overestimated them strongly in the later phases. The largest discrepancy between simulation and measurements was found adjacent to a 13 to 14 days old senescent cluster root, where right from the beginning of the extraction phase the measured citrate-d4 concentration decreased much faster than the simulated one. The cumulative mass recovery for these experiments was 60 to 78 %, except for the cluster root of 13 to 14 days after development: in this experiment,

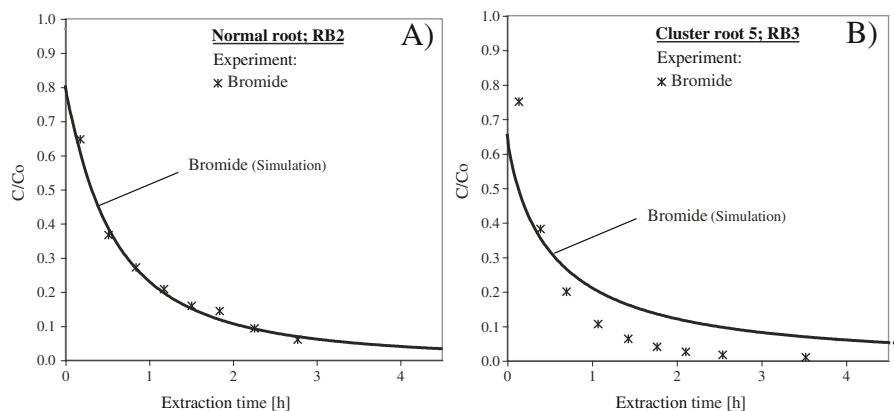


Fig. 2 Extraction breakthrough curves for bromide for two micro push-pull tests adjacent to **a** normal root and **b** cluster root. The root mean squared errors for the best fit were: **a** 0.02; **b** 0.12. C Concentration [mM], C_0 Initial concentration [mM]

Table 1 Overview of the test habitat of the experiments and best-fit results for the bromide data (S , α_L : longitudinal dispersivity, n : porosity) and the citrate-d4 data (k). Corresponding root-mean squared errors (RSME) for the results for bromide are also listed. Reaction rate constants k were calculated assuming first order reaction kinetics

	Test habitat			Best-fit results for bromide/citrate-d4				
	Rhizobox Nr.	Test surrounding	Age of root ^a [days]	Fitting parameters			RMSE	Fitting parameter k [h ⁻¹]
				S	α_L [mm]	n [cm ³ /cm ³]		
S : water saturation; α_L : longitudinal dispersivity; k : reaction rate constant ^a For cluster roots, age is given as days after development ^b No degradation measurable from beginning of extraction phase	RB1	Bulk soil	13–14	0.88	0.40	0.33	0.07	no deg. ^b
	RB1	Normal root		0.90	0.40	0.34	0.03	no deg. ^b
	RB1	Cluster root		0.74	2.30	0.35	0.07	0.61
	RB2	Bulk soil		0.75	0.40	0.32	0.06	no deg. ^b
	RB2	Normal root	3–4	0.84	0.40	0.31	0.02	no deg. ^b
	RB2	Cluster root		0.73	0.40	0.30	0.05	no deg. ^b
	RB2	Cluster root		0.69	0.40	0.35	0.11	–
	RB3	Cluster root 1	3–4	0.91	3.00	0.34	0.02	no deg. ^b
	RB3	Cluster root 2	3–4	0.85	0.40	0.35	0.04	no deg. ^b
	RB3	Cluster root 3	3–4	0.90	3.00	0.35	0.03	no deg. ^b
	RB3	Cluster root 4	4–5	0.79	0.40	0.30	0.04	0.38
	RB3	Normal root	4–5	0.88	0.90	0.35	0.11	no deg. ^b
	RB3	Cluster root 1		0.93	0.90	0.35	0.06	0.71
	RB3	Cluster root 2		0.82	2.40	0.30	0.02	0.51
	RB3	Cluster root 3		0.83	7.60	0.35	0.02	0.50
	RB3	Cluster root 4	5–6	0.83	3.14	0.35	0.18	–
	RB3	Cluster root 5	5–6	0.87	3.13	0.35	0.12	–

only about 27 % of the injected citrate-d4 could be recovered over the extraction time.

Results for the experiments in the surroundings of three cluster roots of different age are shown in Fig. 4 (experiments were repeated after 24 h on the same roots). The divergence between the theoretical simulation for citrate-d4 and the measured concentrations in the experiments with cluster roots after 3 to 4 days of development was again observed only in the rear part of the curves (Fig. 4 a1–3). For the repeated experiments 1 day later, however, the measured citrate-d4 concentrations in the extracted soil solution samples decreased much faster and were from the beginning of the extraction significantly lower compared to the curves predicted by simulation (Fig. 4 b1–3). Consequently, mass recovery was significantly lower for the experiments on the second day: from prior 45–62 % to now 21–29 %. Interpreting the difference between measured and predicted citrate-d4 concentrations being due to microbial degradation, the μ PPT experiments showed only little microbial degradation of citrate-d4 at the beginning of the extraction phase—in bulk soil, adjacent to normal roots and in the surroundings of cluster roots after 3 to

4 days of development, whereas strong degradation took place in the rhizosphere of cluster roots after 4 to 5 days of development. For the latter habitats, first-order degradation rate constants of 0.71, 0.51 and 0.50 h⁻¹ (half-life: $T_{1/2}$ =0.98, 1.36 and 1.37 h) were obtained by fitting the model to the measured data (see Table 1). For the experiment in the surroundings of the senescent cluster root after 13 to 14 days of development, we calculated a degradation rate constant of 0.61 h⁻¹ (half-life: $T_{1/2}$ =1.14 h).

Background citrate

Background citrate concentrations in the sampled soil solutions were determined separately from the injected citrate-d4 concentrations. In bulk soil and in the surroundings of normal roots, we found no measurable citrate concentrations (detection limit: 0.0025 mM) for 4 habitats, while for one habitat the concentrations were <0.02 mM. This is in good agreement with values found by Dessureault-Rompré et al. (2007). Significantly higher concentrations were measured in the rhizosphere of cluster roots (Fig. 4). Citrate is known as the major

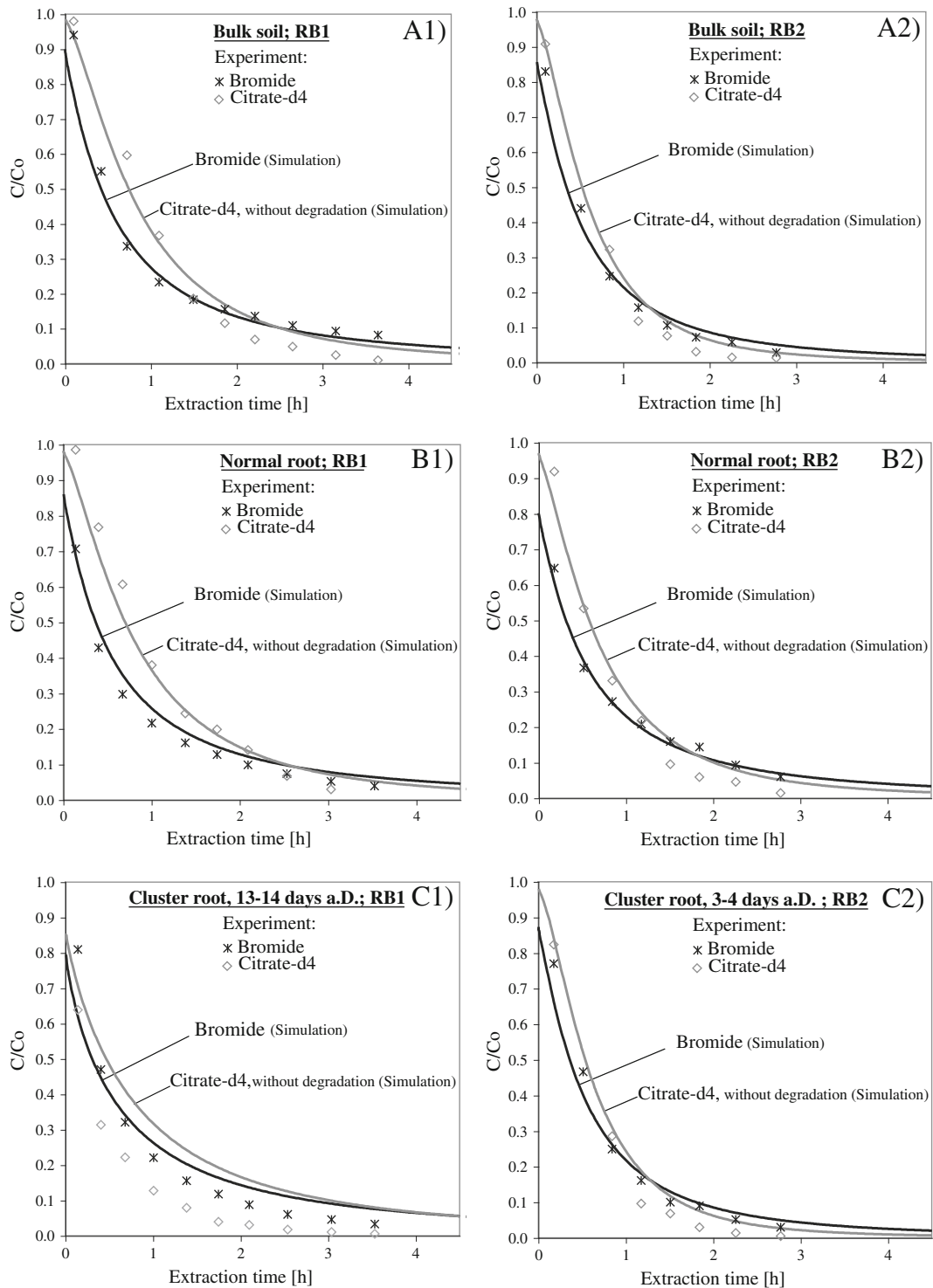


Fig. 3 Extraction-breakthrough curves for bromide and citrate-d4 for two series of micro push-pull test experiments in **a** bulk soil and adjacent to **b** normal roots and **c** cluster roots performed in two different rhizoboxes (*left side*: Rhizobox 1

(RB1); *right side*: Rhizobox 2, RB2). Symbols show the experimental results, *solid lines* show the simulated curves. The development stage of cluster roots is given as days after development (a.D.)

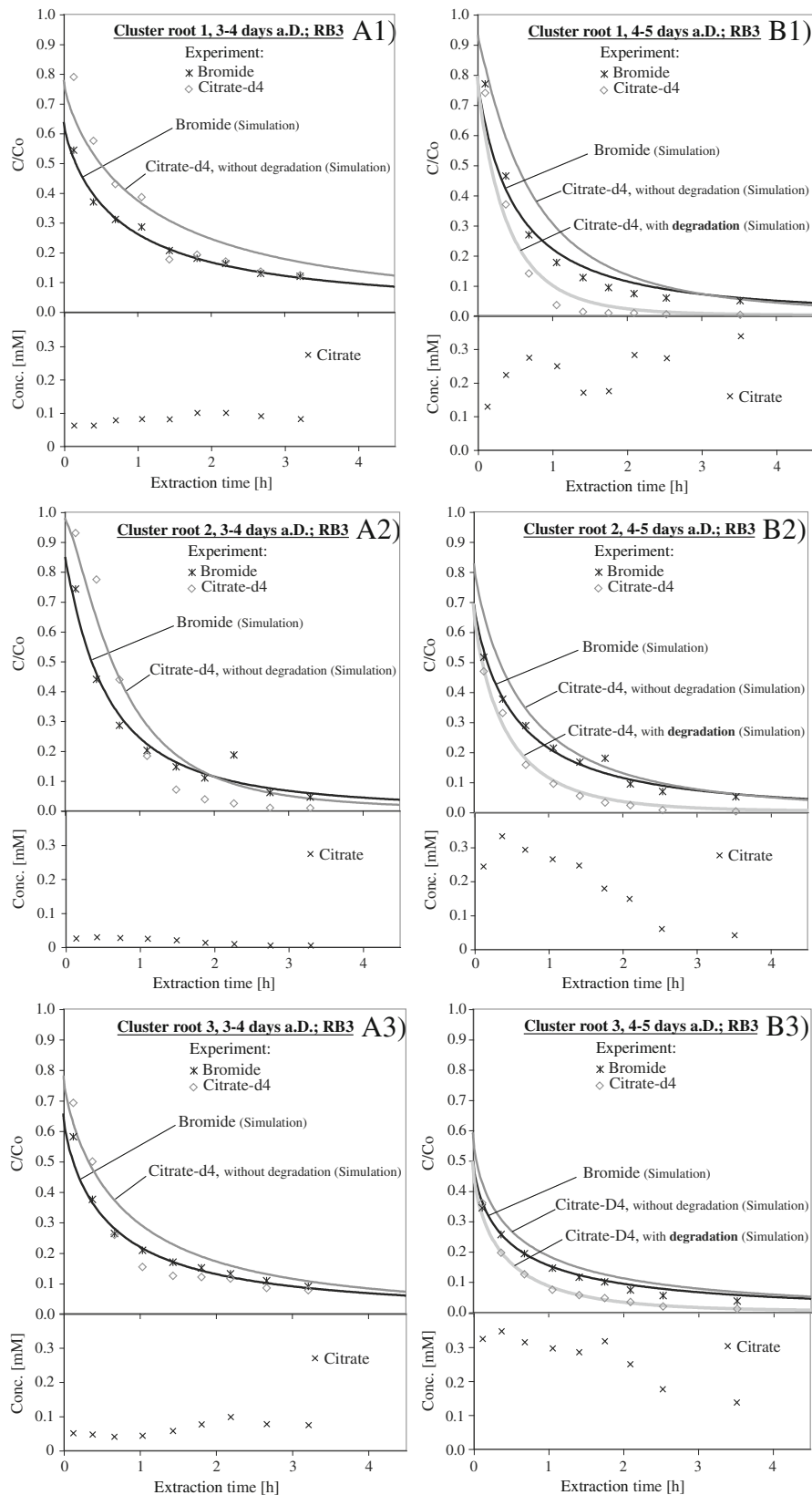


Fig. 4 Extraction-breakthrough curves for μ PPT experiment adjacent to 3 different **cluster** roots (Nr. 1–3) with different development stage of cluster roots of *Lupinus albus*: **a** 3 to 4 days after development and **b** 4 to 5 days after development. Symbols show the experimental results, *solid lines* show the simulated curves. The lower part of the graphics show the measured background citrate concentrations in the extracted soil solution samples

organic anion exuded by *Lupinus albus* (Watt and Evans 1999b). Thus, the increased citrate concentration indicates an exudation activity of the cluster roots—either during the μ PPT experiments or before. Lower concentrations at the beginning of the extraction phases can be due to dilution by the injected test solution, while the later decrease in citrate concentration may be explained by small-scale heterogeneity of background solute concentrations and/or degradation processes. For two habitats, the measurements showed two peaks of citrate concentration, which may indicate a bi-phasic temporal pattern of the exudation activity (Dessureault-Rompré et al. 2007).

In the rhizosphere of the younger cluster roots (3 to 4 days after development), citrate concentrations were rather low and peak values of ≤ 0.1 mM were determined. On the next day, however, citrate concentrations peaked between 0.30 and 0.35 mM in the same habitats. In other words, low concentrations of citrate in the rhizosphere of young cluster roots correlated with no (or only little) biological degradation of citrate-d4 (or at least not at the beginning of extraction), whereas high citrate concentration in the same habitats 1 day later correlated with strong microbial degradation of citrate-d4.

The citrate concentrations measured in the soil solution sampled adjacent to the senescent cluster root 13 to 14 days after development were rather low with a peak value of 0.033 mM, but nevertheless there was a strong immediate degradation of citrate-d4 with a calculated degradation rate constant of 0.61 h^{-1} (half-life: $T_{1/2} = 1.14 \text{ h}$) (Fig. 3c).

The variation in citrate concentration was in accordance with pH values. The pH of the solution samples extracted prior to the tests were significantly higher in the bulk soil and around normal roots than in the rhizosphere of active cluster roots: $\text{pH}_{\text{bulk soil}; (n=6)} = 6.0$; $\text{pH}_{\text{normal root}; (n=5)} = 5.8$; $\text{pH}_{\text{cluster root, age: 3–4 days}; (n=4)} = 4.6$; $\text{pH}_{\text{cluster root, age 5–6 days}; (n=2)} = 3.5$; $\text{pH}_{\text{cluster root, age 13–14 days}; (n=1)} = 6.0$.

Discussion

Zone of influence

The zone of influence is of high relevance for the μ PPT method, since the purpose of the test is to assess rhizosphere processes on a small scale with minimal impact on the soil. The impacted soil volume of a μ PPT is primarily determined by the injection volume and the radius r of the impacted soil cylinder around the suction cup may be computed using (Knecht et al. 2013)

$$r = \sqrt{\frac{Qt}{\pi b \theta} + r_{sc}^2} \quad (1)$$

where Q is the injection rate, t is the injection time, b is the length of the suction cup, θ is the water content and r_{sc} is the radius of the suction cup. With decreasing water saturation the impacted soil volume increases accordingly. For a saturation of 0.7 for example, the impacted sand pack volume would have been approximately 490 to 660 μl and the radius of the corresponding cylinder around the suction cup 4 to 4.6 mm with no mixing. Considering also diffusion/dispersion, the soil volume that actually came into contact with the injected solution was even larger, depending on the extent of these processes. In this case, the calculated radius can be understood as the location where $C/C_o = 0.5$ for a non-sorbing, non-reactive tracer.

For an experiment with an extraction rate of 2.3 to 2.6 $\mu\text{l}/\text{min}$ and an extraction time of 2.5 h, the sampled soil volume would be about 1,400 to 1,600 μl (equivalent to a soil cylinder around the suction cup with a radius of 6.7 to 7.1 mm) for a water saturation of 0.7. The width of the rhizosphere can vary considerably around a root, depending on the process of interest and can extend from <1 mm (e.g. for poorly mobile nutrients like phosphate) up to several centimeters (e.g. for mobile nutrients like nitrate) (cf. Darrah 1993; Hinsinger et al. 2005). Thus, the μ PPT method shows potential for studying rhizosphere processes in situ. The substitution of micro-suction cups by micro capillaries, as used by Tandy et al. (2013) for collection of soil solution might allow further miniaturization of the μ PPT. Increased spatial resolution could then be achieved by placing several micro-suction cups (or micro capillaries) around the same root. The micro capillaries could substitute the micro-suction cups used, but not the μ PPT, which is an experimental method and not a device.

Hydraulic characterization of porous medium

To adapt the μ PPT method to a porous medium populated with roots and associated rhizosphere organisms, it was essential to accommodate changes in its hydraulic properties due to microorganisms and root growth. Information about the impact of the colonization of our sand packs by roots and microbes on flow and transport characteristics was gained from the breakthrough curve of the simultaneously injected non-reactive tracer. By varying three relevant hydraulic parameters (porosity, n ; longitudinal dispersivity, α_L ; initial hydraulic pressure, P_i) in the numerical simulation during fitting, we forced the simulated curves to match the respective measured data for bromide concentrations adjacent to roots (normal roots and cluster roots) and in bulk soil as close as possible. Our results showed that the optimized values of this parameter set allowed for most experiments to adequately describe the measured bromide breakthrough curves and thus for adapting the μ PPT method to different habitats. Nonetheless, due to numerous assumptions made in the simulations (Knecht et al. 2013), the best-fit parameters may not reflect actual physical conditions of the porous medium. While fitting parameters P_i and n affect the shape of the breakthrough curves in a similar fashion, α_L affects their shape in a different way.

From our results, we can see a tendency to higher dispersivity values in the rhizosphere of cluster roots than around normal roots and in bulk soil. This might be due to more heterogeneity caused by enhanced activity of roots or microorganism or quite simply because of the differences in root architecture. Cluster roots form dense rootlets with lengths of up to 9 mm in our experiments, which is in the range given by Dinkelaker et al. (1995). These differences to bulk soil and around roots may be sufficient to explain the different overall flow/transport behavior of solutes in μ PPT adjacent to cluster roots.

In four experiments, the RMSE from the simulations was rather high compared to the other tests, and the model could not be fitted to the measured data within the range of predetermined physical model parameters. In these experiments, a faster decrease of the measured bromide concentration was observed than was predicted, and mass recovery was considerably lower than for the other experiments. A plausible explanation for this failure might be preferential flow. Under unsaturated conditions, preferential flow might become a limiting phenomenon for the μ PPT as the flow patterns between

injection and extraction phase will not be entirely reversible any more. Preferential flow can result from macro-pore flow through e.g. cracks and fissures, where during injection the test solution rapidly flows away from the injection point and cannot be recovered during the extraction phase. Such preferential flow was observed in preliminary μ PPT with Acid Red 1 as tracer performed in one of the rhizoboxes with plants. The formation of cracks and fissures can be induced especially in the surroundings of roots by a high capillary water tension associated with root water uptake (Pierret et al. 2007). However, note that high mass recovery of solutes during a PPT is not a prerequisite for accurate rate determination (Haggerty et al. 1998).

Bromide has often been used as a conservative tracer in subsurface solute transport studies due to its generally low background concentration and low reactivity in soils (Schnabel et al. 1995). If bromide comes into contact with roots, also uptake by plants can be an important transport pathway (Kohler et al. 2005). The fraction of bromide taken up by plants such as wheat, potato and ryegrass in field tracer experiments was found to range from <2 % (Gish and Jury 1982) to 53 % (Kung 1990) of the applied tracer mass, depending on factors such as rate of tracer application, soil drainage, growth season and nitrogen fertilizer application. While plants were exposed to the tracer for several weeks to months in these studies, our tests lasted only 4 to 4.5 h. In addition, by covering the plants with a plastic bag we minimized transpiration-induced root water extraction during the tests. Furthermore, we used chloride, which is chemically analogous to bromide, as background electrolyte in all our experiments. As a competing anion, chloride reduces the uptake of bromide by plants (cf. also Epstein 1953; Sheppard and Evenden 1992). Taken all these aspects together, bromide uptake by the experimental plants was considered negligible in our tests. Also, the relative mass recovery of bromide from rhizosphere soil compared to root-free bulk soil indicated no significant uptake.

Microbial degradation of citrate-d4

The citrate concentrations in the rhizosphere suggest that the young cluster roots were at the beginning of the exudation activity. According to Watt and Evans (1999b), the efflux of citrate is linked among others to proteoid root development and they reported a beginning of citrate efflux by *Lupinus albus* after about 4 days

of development with a diurnal pulse over 2 or 3 days. Moreover, Shane et al. (2003) observed highest citrate exudation rates for cluster roots after 3 to 5 days of development. Our results from the citrate measurements fit well into these observations, although these studies were done in hydroponic systems and roots can behave quite differently in different growth media or in soils (Marschner 1995; Watt and Evans 1999a). Dessureault-Rompré et al. (2006) observed a similar behavior of citrate exudation in experiments with *Lupinus albus* grown in soil. The same authors reported a high temporal variability in the beginning of citrate exudation between individual cluster roots in a subsequent paper (Dessureault-Rompré et al. 2007). In line with our observations of lower pH around the active cluster roots, Dessureault-Rompré et al. (2006) reported a decrease in soil pH adjacent to cluster roots from 6.9 to 5.4 during citrate exudation. The mechanism for organic acid efflux by roots is not fully clarified. There is evidence that citrate is excreted concomitantly with protons (Watt and Evans 1999a), although Jones (1998) suggested that protons and organic acid are exuded separately of each other, but in spatially coordinated transport events.

The breakthrough curves of the extracted citrate-d4 varied considerably between tests performed adjacent to normal roots, in bulk soil and in the rhizosphere of cluster roots at different stages of development, with significantly lower citrate-d4 concentrations in the rhizosphere of active cluster roots and of a senescent cluster root. The fast initial decrease of citrate-d4 concentrations at the beginning of the extraction phase could not be predicted by the model (using the same flow/transport conditions as for bromide, except for the diffusion coefficient). The assumption that biological degradation is responsible for this discrepancy between measured citrate-d4 concentrations compared to the theoretical simulations is based on the following points: (1) In a previous study, we found no measurable adsorption of citrate to the sand used also in the present study, and citrate behaved like a non-reactive tracer in μ PPT under sterile conditions (Knecht et al. 2011). (2) We assume that citrate-d4 has the same adsorption behavior as citrate. (3) Root activity and microbial processes (e.g. substrate consumption, biomass growth, respiration, production of extracellular cells, release of gaseous byproducts, attachment/detachment to and from surfaces etc.) can change physical and chemical properties of a porous medium (cf. e.g. Baveye et al. 1998; Frankenberger et al. 1979; Hinsinger et al. 2003;

Rockhold et al. 2004). Thus effects on the sorption of the solute cannot be ruled out completely under non-sterile conditions. Different parameters (e.g. adsorption parameters, degradation rate) have a very specific effect on the shape of a breakthrough curve (Schroth et al. 2001). Thus, the numerical simulation is very important and helps to distinguish between different processes. Simulations in which varying degrees of sorption (different K_D values; K_D : solid-aqueous partition coefficient) were assumed to influence the extraction of citrate show that the concentration of citrate would be decreased significantly in the beginning of extraction phase and increased slightly at later stage of extraction phase (compared to not adsorbed solutes), resulting in a completely different form of the predicted citrate concentrations (see Fig. 5). 4) There is good agreement between data and model in Fig. 4b using a reaction term in the simulations, whereas the data points may not be fitted using a sorption term.

The measured concentrations of citrate-d4 were lower than predicted towards the end of all these tests. We assume that this indicates microbial degradation of citrate-d4 after a lag-phase, thus 1 to 1.5 h after beginning the extraction (about 2.5 h after the start of the experiment). The assumption of a delayed start of microbial degradation is supported by an incubation experiment with citrate and soil containing soil microflora

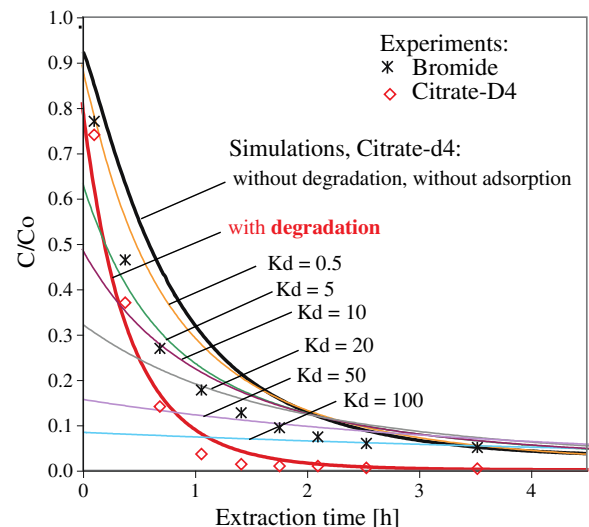


Fig. 5 Extraction-breakthrough curve for bromide and citrate-d4 with simulations including degradation of citrate and sorption of citrate to sand at different K_D values (solid-aqueous partition coefficient). Symbols show the experimental results, solid lines show the simulated curves. The experimental data correspond to the results from the experiment shown in Fig. 4 B1

(Weisskopf et al. (2006). These authors found that citrate degradation was negligible during the first 3 h of incubation and then rapidly accelerated to be complete after 24 h.

The experiments performed on the same cluster roots 1 day later showed significantly lower citrate-d4 concentrations at the beginning of the extraction phase compared to the simulation and the measured background citrate concentrations were much higher than the day before with peak values between 0.3 and 0.35 mM. We conclude that at the time when we injected our test solution on the second day, the micro-organisms were already stimulated by release of citrate from the roots. In addition, degradation on the second day (adjacent to cluster roots after 4 to 5 days of development) may have been enhanced by induction of microbial activity by the tests conducted at the same location the day before. But strong degradation and a peak value for background citrate of 0.33 mM were also found adjacent to 4 to 5 days old cluster roots, which had not been tested the day before.

From the second-day tests, we determined first-order degradation rate constants for citrate-d4 that were in the range of 0.38 to 0.71 h⁻¹ in the rhizosphere of active cluster roots after 4 to 5 days of development (corresponding half-life: $T_{1/2}$ =1.81 to 0.98 h). From the results of the experiment in the rhizosphere of the senescent cluster root, we determined a degradation rate of 0.57 h⁻¹ and measured only low concentrations of background citrate. Jones et al. (2003) determined half-lives of organic acids turnover ranging from 1 to 5 h in organic topsoils and from 5 to 12 h in subsoils. Van Hees et al. (2003) determined a mean half-life of 2 to 6 h for citrate which was incubated in forest soils from organic surface horizons. Jones (1998) estimated that decomposition rates in rhizosphere soil were 2–3 fold higher than in bulk soil. On the other hand, Weisskopf et al. (2006) revealed complex strategies by which *Lupinus albus* protects secreted organic anions from microbial degradation. As far as we know, this was the first time that in situ rates have been determined for the degradation of organic acids in the rhizosphere.

The rhizosphere is often described as a “hot spot” of soil biological activity, which is attributed to the release of exudates by roots that serve as food for rhizosphere bacteria, and also act as chemo-attractants inducing microbes to move towards the roots (e.g. Jones 1998). This study underlines the ecological role of the rhizosphere as a special soil micro-habitat with high

degradation potentials and high turnover rates. Our results show that the μ PPT has the potential to become a useful technique to study rhizosphere processes in situ and—because it is a non-destructive method with only minimal disturbance—to monitor them also over time. Further research is required to adapt the method to more complex systems (e.g. well-graded soil etc.) with e.g. more small-scale heterogeneity and possible influences of solute adsorption have to be taken into account. Our results complement previous findings based on modeling (Knecht et al. 2013), which demonstrated the feasibility of μ PPT in soils of varying soil physical and hydraulic properties.

Acknowledgments The authors thank M. Oostrom and M.D. White (Pacific Northwest National Laboratory) for providing STOMP and T. Zwickl (Wädenswil, Switzerland) for Ruby scripting of STOMP fitting and scanning methods. The authors also thank M. Keller (ETHZ) for assistance in designing and constructing the equipment, René Schönenberger (Eawag, Switzerland) for analysis of citrate-d4 and citrate and Jörg Luster (WSL, Switzerland) for providing a rhizobox prototype and the ceramic capillaries for the micro-suction cups. This research was funded by the Swiss National Science Foundation projects 200021–113792 and 200020–127139.

References

- Baveye P, Vandevivere P, Hoyle BL, DeLeo PC, de Lozada DS (1998) Environmental impact and mechanisms of the biological clogging of saturated soils and aquifer materials. *Crit Rev Environ Sci Technol* 28:123–191
- Blossfeld S, Gansert D, Br T, Kuhn AJ, Lösch R (2011) The dynamics of oxygen concentration, pH value, and organic acids in the rhizosphere of *Juncus* spp. *Soil Biol Biochem* 43: 1186–1197
- Buffle J, Zhang Z, Startchev K (2007) Metal flux and dynamic speciation at (Bio)interfaces. part 1: critical evaluation and compilation of physicochemical parameters for complexes with simple Ligands and Fulvic/Humic substances. *Environ Sci Technol* 41:7609–7620
- Darrah PR (1993) The rhizosphere and plant nutrition—a quantitative approach. *Plant Soil* 156:1–20
- Dessureault-Rompré J, Nowack B, Schulin R, Luster J (2006) Modified micro suction cup/rhizobox approach for the in-situ detection of organic acids in rhizosphere soil solution. *Plant Soil* 286:99–107
- Dessureault-Rompré J, Nowack B, Schulin R, Luster J (2007) Spatial and temporal variation in organic acid anion exudation and nutrient anion uptake in the rhizosphere of *Lupinus albus* L. *Plant Soil* 301:123–134
- Dieffenbach A, Göttlein A, Matzner E (1997) In-situ soil solution chemistry in an acid forest soil as influenced by growing

- roots of Norway spruce (*Picea abies* [L.] Karst). Plant Soil 192:57–61
- Dinkelaker B, Hengeler C, Marschner H (1995) Distribution and function of proteoid roots and other root clusters. Bot Acta 108:183–200
- Drever JI, McKee CR (1980) The push-pull test—A method for evaluating formation adsorption parameters for predicting the environmental effects on in-situ coal gasification and uranium recovery. In Situ 4:181–206
- Epstein E (1953) Mechanism of ion absorption by roots. Nature 171:83–84
- Frankenberger WT, Troeh FR, Dumenil LC (1979) Bacterial effects on hydraulic conductivity of soils. Soil Sci Soc Am J 43:333–338
- Gish TJ, Jury WA (1982) Estimating solute travel-times through a crop root zone. Soil Sci 133:124–130
- Göttlein A, Hell U, Blasek R (1996) A system for microscale tensiometry and lysimetry. Geoderma 69:147–156
- Göttlein A, Heim A, Matzner E (1999) Mobilization of aluminium in the rhizosphere soil solution of growing tree roots in an acidic soil. Plant Soil 211:41–49
- Gregory PJ, Hinsinger P (1999) New approaches to studying chemical and physical changes in the rhizosphere: an overview. Plant Soil 211:1–9
- Haggerty R, Schroth MH, Istok JD (1998) Simplified method of “push-pull” test data analysis for determining in situ reaction rate coefficients. Ground Water 36:314–324
- Hinsinger P, Plassard C, Tang CX, Jaillard B (2003) Origins of root-mediated pH changes in the rhizosphere and their responses to environmental constraints: a review. Plant Soil 248:43–59
- Hinsinger P, Gobran GR, Gregory PJ, Wenzel WW (2005) Rhizosphere geometry and heterogeneity arising from root-mediated physical and chemical processes. New Phytol 168:293–303
- Hinsinger P, Bengough AG, Vetterlein D, Young IM (2009) Rhizosphere: biophysics, biogeochemistry and ecological relevance. Plant Soil 321:117–152
- Istok JD, Humphrey MD, Schroth MH, Hyman MR, Oreilly KT (1997) Single-well, “push-pull” test for in situ determination of microbial activities. Ground Water 35:619–631
- Johnson JF, Allan DL, Vance CP, Weiblen G (1996) Root carbon dioxide fixation by phosphorus-deficient *Lupinus albus*—contribution to organic acid exudation by proteoid roots. Plant Physiol 112:19–30
- Jones DL (1998) Organic acids in the rhizosphere—a critical review. Plant Soil 205:25–44
- Jones DL, Dennis PG, Owen AG, van Hees PAW (2003) Organic acid behavior in soils—misconceptions and knowledge gaps. Plant Soil 248:31–41
- Keerthisinghe G, Hocking PJ, Ryan PR, Delhaize E (1998) Effect of phosphorus supply on the formation and function of proteoid roots of white lupin (*Lupinus albus* L.). Plant Cell Environ 21:467–478
- Knecht K, Schroth MH, Schulin R, Nowack B (2011) Development and evaluation of micro push-pull tests to investigate micro-scale processes in porous media. Environ Sci Technol 45:6460–6467
- Knecht K, Nowack B, Schulin R, Schroth MH (2013) Micro push-pull tests for investigation of small-scale processes in unsaturated porous media. Vadose Zone J 12. doi:10.2136/vzj2012.0087.
- Kohler A, Abbaspour KC, Fritsch M, Schulin R (2005) Solute recycling by crops and leaching in a drained arable soil. Eur J Soil Sci 56:145–153
- Kung KJS (1990) Influence of plant uptake on the performance of bromide tracer. Soil Sci Soc Am J 54:975–979
- Liang RX, Li CJ (2003) Differences in cluster-root formation and carboxylate exudation in *Lupinus albus* L. under different nutrient deficiencies. Plant Soil 248:221–227
- Luster J, Göttlein A, Nowack B, Sarret G (2009) Sampling, defining, characterising and modeling the rhizosphere—the soil science tool box. Plant Soil 321:457–482
- Marschner H (1995) Mineral nutrition of higher plants. Academic Press, London
- Neumann G, Römhild V (1999) Root excretion of carboxylic acids and protons in phosphorus-deficient plants. Plant Soil 211:121–130
- Neumann G, Massonneau A, Martinoia E, Römhild V (1999) Physiological adaptations to phosphorus deficiency during proteoid root development in white lupin. Planta 208:373–382
- Pierret A, Doussan C, Capowiez Y, Bastardie F, Pages L (2007) Root functional architecture: a framework for modeling the interplay between roots and soil. Vadose Zone J 6:269–281
- Puschenreiter M, Wenzel WW, Wieshammer G, Fitz WJ, Wiczorek S, Kanitsar K, Kollensperger G (2005) Novel micro-suction-cup design for sampling soil solution at defined distances from roots. J Plant Nutr Soil Sci 168:386–391
- Rockhold ML, Yarwood RR, Selker JS (2004) Coupled microbial and transport processes in soils. Vadose Zone J 3:368–383
- Schnabel RR, Stout WL, Shaffer JA (1995) Uptake of hydrologic tracer (bromide) by ryegrass from well and poorly-drained soils. J Environ Qual 24:888–892
- Schreiber C, Zeng B, Temperton V, Rascher U, Kazda M, Schurr U, Höltkemeier A, Kuhn A (2011) Dynamics of organic acid occurrence under flooding stress in the rhizosphere of three plant species from the water fluctuation zone of the Three Gorges Reservoir, P.R. China. Plant Soil 344:111–129
- Schroth MH, Ahearn SJ, Selker JS, Istok JD (1996) Characterization of miller-similar silica sands for laboratory hydrologic studies. Soil Sci Soc Am J 60:1331–1339
- Schroth MH, Istok JD, Haggerty R (2001) In situ evaluation of solute retardation using single-well push-pull tests. Adv Water Resour 24:105–117
- Shane MW, De Vos M, De Roock S, Lambers H (2003) Shoot P status regulates cluster-root growth and citrate exudation in *Lupinus albus* grown with a divided root system. Plant Cell Environ 26:265–273
- Sheppard SC, Evenden WG (1992) Response of some vegetable crops to soil-applied halides. Can J Soil Sci 72:555–567
- Simunek J, Jarvis NJ, van Genuchten MT, Gardenas A (2003) Review and comparison of models for describing non-equilibrium and preferential flow and transport in the vadose zone. J Hydrol 272:14–35
- Tandy S, Schulin R, Nowack B (2006) The influence of EDDS on the uptake of heavy metals in hydroponically grown sunflowers. Chemosphere 62:1454–1463
- Tandy S, Brittain S, Grail B, McLeod C, Paterson E, Tomos AD (2013) Fine scale measurement and mapping of uranium in soil solution in soil and plant-soil microcosms, with special

- reference to depleted uranium. *Plant Soil*. doi:[10.1007/s11104-11012-11526-11107](https://doi.org/10.1007/s11104-11012-11526-11107)
- Tidwell VC, Glass RJ (1994) X-ray and visible-light transmission for laboratory measurement of 2-dimensional saturation fields in thin-slab systems. *Water Resour Res* 30:2873–2882
- van Hees PAW, Jones DL, Godbold DL (2003) Biodegradation of low molecular weight organic acids in a limed forest soil. *Water Air Soil Pollut Focus* 3:121–144
- Vetterlein D, Jahn R (2004) Gradients in soil solution composition between bulk soil and rhizosphere—in situ measurement with changing soil water content. *Plant Soil* 258:307–317
- Vetterlein D, Marschner H, Horn R (1993) Microtensiometer technique for insitu measurement of soil matric potential and root water extraction from a sandy soil. *Plant Soil* 149: 263–273
- Watt M, Evans JR (1999a) Linking development and determinacy with organic acid efflux from proteoid roots of white lupin grown with low phosphorus and ambient or elevated atmospheric CO₂ concentration. *Plant Physiol* 120:705–716
- Watt M, Evans JR (1999b) Proteoid roots. Physiology and development. *Plant Physiol* 121:317–323
- Weisskopf L, Abou-Mansour E, Fromin N, Tomasi N, Santelia D, Edelkott I, Neumann G, Aragno M, Tabacchi R, Martinoia E (2006) White lupin has developed a complex strategy to limit microbial degradation of secreted citrate required for phosphate acquisition. *Plant Cell Environ* 29:919–927
- White MD, Oostrom M (2000) STOMP. Subsurface transport over multiple phases, version 2.0, Theory Guide PNNL-12030. Pacific Northwest National Laboratory, Richland