⁹Field Science Center for Northern Biosphere, Hokkaido University, Sapporo, Japan

¹¹CREAF, Campus Universitat Autònoma de Barcelona, E08193, Bellaterra, Spain

¹⁰ILTER Information Management Committee, Jerusalem, Israel

30 31

- ¹²Department of Botany and Biodiversity Research, Division of Conservation Biology, Vegetation 33 34 Ecology and Landscape Ecology, University of Vienna, Rennweg 14, 1030 Vienna Austria
- 35 ¹³Agrosphere Institute (IBG-3), Forschungszentrum Jülich, 52425 Jülich, Germany
- 36 ¹⁴Cooperative Institute for Research in Environmental Sciences, University of Colorado Boulder, 37 Boulder, CO, USA and NOAA Chemical Science Laboratory, Boulder, CO, USA
- 38 ¹⁵US Forest Service, Northern Research Station, Durham, New Hampshire, USA
- 39 ¹⁶Environment Agency Austria, A-1090 Vienna, Austria
- 40 ¹⁷Division of Biology, Kansas State University, Manhattan, Kansas, USA
- 41 ¹⁸NIBIO, Norwegian Institute of Bioeconomy Research, Ås, Norway
- 42 ¹⁹Institute of Biology, University of Latvia, Jelgavas iela 1, Riga, LV-1004, Latvia and Latvian 43 Environment, Geology and Meteorology Center, Maskavas iela 165, LV-1019, Riga, Latvia 44
 - ²⁰Department of Crop and Soil Sciences, Oregon State University, Corvallis OR, USA
- 45 ²¹Graduate Institute of Bioresources, National Pingtung University of Science and Technology, 1 46 Shuefu Rd, Neipu, Pingtung 912, TAIWAN
 - ²²Dept. Forest Ecology and Management, Swedish University of Agricultural Sciences, 901 83 Umeå, Sweden
- 49 ²³Department of Life Science, National Taiwan Normal University, Taipei, Taiwan
- 50 ²⁴Center for Nuclear Energy in Agriculture, University of São Paulo, Piracicaba, Brazil.
- ²⁵Northwest German Forest Research Institute, Grätzelstr. 2, 37079 Göttingen, Germany 51
- 52 ²⁶Centre for Ecology, Evolution and Environmental Changes, Universidade de Lisboa, Lisbon, 53 Portugal
- 54 ²⁷CNR Water Research Institute, L.go Tonolli 50, I 28922 Verbania Pallanza, Italy
- 55 ²⁸Swiss Federal Institute for Forest, Snow and Landscape Research, Zürcherstr. 111, CH-8903 56 Birmensdorf, Switzerland
- 57 ²⁹Department of Ecology, University of Innsbruck, 6020 Innsbruck, Austria
- 58 ³⁰Integrated Monitoring Department, Institute for Ecology of Industrial Area, Katowice, Poland
- 59 ³¹Water and Environmental Engineering Research Group, Aalto University, P.O. Box 15200, FI-60 00076 Aalto, Espoo, Finland

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- 63 Keywords: bulk nitrogen deposition, LTER, atmospheric pollution, throughfall, watershed, water
- 64 quality

Abstract

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Previous studies have evaluated how changes in atmospheric nitrogen (N) inputs and climate affect stream N concentrations and fluxes, but none have synthesized data from sites around the globe. We identified variables controlling stream inorganic N concentrations and fluxes, and how they have changed, by synthesizing 20 time series ranging from 5 to 51 years of data collected from forest and grassland dominated watersheds across Europe, North America, and East Asia and across four climate types (tropical, temperate, Mediterranean, and boreal) using the International Long-Term Ecological Research Network. We hypothesized that sites with greater atmospheric N deposition have greater stream N export rates, but that climate has taken a stronger role as atmospheric deposition declines in many regions of the globe. We found declining trends in bulk ammonium and nitrate deposition, especially in the longest time-series, with ammonium contributing relatively more to atmospheric N deposition over time. Among sites, there were statistically significant positive relationships between (1) annual rates of precipitation and stream ammonium and nitrate fluxes and (2) annual rates of atmospheric N inputs and stream nitrate concentrations and fluxes. There were no significant relationships between air temperature and stream N export. Our long-term data shows that although N deposition is declining over time, atmospheric N inputs and precipitation remain important predictors for inorganic N exported from forested and grassland watersheds. Overall, we also demonstrate that long-term monitoring provides understanding of ecosystems and biogeochemical cycling that would not be possible with short-term studies alone.

Introduction

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Nitrogen (N) is an essential, but often limiting element for net primary production in temperate terrestrial ecosystems (LeBauer and Treseder 2008), while phosphorus (P) availability often controls patterns of primary productivity in tropical terrestrial ecosystems (Condit et al. 2013; Turner et al. 2018). Both N and P are vital nutrients in aquatic ecosystems as well (Wurtsbaugh et al. 2019). Studies published over the last several decades have aimed to better understand the effects of human activities on the biogeochemical cycling of N and the consequences of elevated rates of N atmospheric deposition for N export to streams (e.g., Dise and Wright 1995; Shibata et al. 2001; Aber et al. 2003; Dise et al. 2009). For example, human activities, including combustion of fossil fuels, planting crops associated with N fixing bacteria, and synthesis and application of fertilizers have led to elevated rates of N fixation and atmospheric deposition around the globe (Galloway et al. 2008). Even relatively low rates of atmospheric inputs of N can fertilize plants and stimulate growth and carbon sequestration (Thomas et al. 2010; Yu et al. 2014; Etzold et al. 2020). However, high rates of N inputs can saturate biological demand for N by plants and microbes, as well as soil exchange sites, leading to a series of negative impacts known as N saturation (Aber et al. 1989; 1998) or the N cascade (Galloway et al. 2003). These negative effects include nutrient imbalances in plants, acidification of nearby waterways, and eutrophication of lakes and coastal areas (Galloway et al. 2008). Even though a large proportion (up to 90%) of atmospheric N inputs is retained within terrestrial ecosystems, particularly in watersheds with relatively low levels of atmospheric N inputs (<10 kg N ha⁻¹ yr⁻¹), similar amounts of N can move into terrestrial ecosystem sinks (e.g., plants and soils) as are lost in leachate and these processes can occur simultaneously (Lovett and Goodale 2011).

Although current rates of atmospheric N deposition are elevated compared to pre-industrial levels in many parts of the globe (Galloway et al. 2008), N deposition, especially as nitrate, has declined over the last few decades across much of North America (Templer et al. 2012; Lloret and Valiela 2016) and Europe (Waldner et al. 2014; Theobald et al. 2019), and more recently in parts of China (Wen et al. 2020; Zhao et al. 2021). Recent declines in nitrate deposition have been attributed to government policies that limit emissions of N oxides, but rates of atmospheric N deposition remain high and are increasing in parts of Asia (Ge et al. 2020; Takahashi et al. 2020; Wen et al. 2020). Further, ammonium deposition is still elevated and contributes an increasing proportion of total N deposition in North America (Templer et al. 2012; Li et al. 2016) and Europe (EMEP Status Report 2021).

A previous meta-analysis of input-output data and N addition experiments across Europe (Dise and Wright 1995) demonstrated a significant positive relationship between atmospheric N deposition in throughfall and N flux in streams. At N deposition rates below 8 kg N ha⁻¹ yr⁻¹, watersheds could effectively remove most of the N inputs, and at N deposition rates above about 25 kg N ha⁻¹ yr⁻¹, significant N leaching occurred (Dise and Wright 1995). Between N deposition rates of 8 and 25 kg N ha⁻¹ yr⁻¹, retention of atmospheric N inputs varied widely across the forested watersheds studied with sites that have organic soil C:N ratios less than 25 having significantly greater N leaching rates than sites with higher C:N ratios (MacDonald et al. 2002). Other studies also showed increasing N losses to streams with increasing deposition above a similar critical threshold (Shibata et al. 2001; Aber et al. 2003; Bernot and Dodds 2005; Mulholland et al. 2009; Pardo et al. 2011; Nishina et al. 2017; Sugimoto and Tsuboi 2017; Vuorenmaa et al. 2018).

In addition to changes in atmospheric N deposition, shifts in climatic conditions can have important implications for N cycling in terrestrial ecosystems and have consequences for the export

of N into nearby waterways (Greaver et al. 2016). Changes in the hydrologic cycle can influence N sinks in terrestrial watersheds, in turn affecting N cycling processes and the potential movement of N from terrestrial to aquatic ecosystems. With warmer temperatures, runoff decreases due to higher evapotranspiration, but the hydrologic cycle can also be intensified with more extreme rainfall, higher annual precipitation and drought events, as well as with reductions in snowpack depth and duration (Park et al. 2010; Dirnböck et al. 2020). Greater rates of precipitation and more frequent high-intensity events can flush out N from soils, leading to greater stream N export (Whitehead et al. 2009). Droughts followed by wet events can also lead to higher N concentrations in streams (Whitehead et al. 2009; Leitner et al. 2020), potentially leading to severe water quality problems (Loecke et al. 2017).

Many other factors, however, are important in determining the pathways of N across the soilplant continuum and thus, altering N losses to nearby aquatic ecosystems. For example, changes in
climate can affect coupled plant-soil N dynamics within terrestrial ecosystems that could lead to
changing patterns of N loss to nearby waterways. Further, N demand by terrestrial vegetation is
increasing over time relative to availability, in part due to rising temperatures, lengthening growing
seasons, and increasing atmospheric concentrations of carbon dioxide (CO₂; Craine et al. 2018). In
contrast, trees subjected to increased CO₂ concentrations may increase N availability in the soil,
presumably through a priming effect (Schleppi et al. 2019). In-stream processing can remove
significant amounts of N loading, depending on watershed vegetation, hydrology, and woody debris
characteristics (Bernhardt et al. 2005, Adams et al. 2014). Similarly, forest growth, and loading of
coarse woody debris, can significantly increase N retention in forested landscapes compared to
atmospheric loading (Lajtha 2020). Moreover, accumulated deposition of N can affect the response

of forest growth to climatic changes both positively (Dirnböck et al. 2017), but also negatively in case of other nutrient limitations (Braun et al. 2017).

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To our knowledge, a global synthesis of studies evaluating the role of atmospheric N deposition and climate on stream inorganic N concentrations and fluxes from natural ecosystems has not yet been conducted. We used datasets from sites across the International Long-Term Ecological Research Network (ILTER) to address this gap in knowledge. Long-term monitoring of specific ecosystems can reveal patterns that would not otherwise be apparent from short-term studies (Hobbie et al. 2003; Lovett et al. 2007; Shibata et al. 2015; Mirtl et al. 2018). Several countries around the globe have long-term monitoring networks and the ILTER network was established in 1993 to connect these national networks, including sites in 44 countries. The ILTER network focuses primarily on ecosystem structure, function, and services (Dirnböck et al. 2019; Mirtl et al. 2018). For example, researchers have used the ILTER network to evaluate changes in atmospheric CO₂ concentrations (Curcoll et al. 2019), linkages between species composition of macrophytes and rates of primary production (Germ et al. 2019), and connections between past climate change and plant and soil composition (Figueroa-Rangel 2019). Additional ILTER-related research includes evaluations of human impacts on coastal and ocean ecosystems (Muelbert et al. 2019), temporal changes in biodiversity in response to regional climate and local conditions (Pilotto et al. 2020), and impacts of climate and atmospheric N deposition on rates of litter decomposition (Kwon et al. 2021).

We examined trends over time in atmospheric N deposition, climate, and stream inorganic N concentrations and fluxes from non-urban and non-agricultural ecosystems across sites in the ILTER network (Figure 1). Our objective was to identify variables controlling stream N concentrations and fluxes and to determine how these drivers change over time. We hypothesized that sites with greater rates of atmospheric N deposition have greater concentrations and fluxes of stream N. However, with

declining rates of atmospheric N deposition in many regions, we expected climate to be a stronger driver than atmospheric N deposition on stream N concentrations and fluxes.

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Methods

We assembled data from unmanaged (i.e., non-urban and non-agricultural) forest and grassland ecosystems in ILTER sites around the globe (Figure 1; Table 1; Templer et al. 2022). We contacted all ILTER sites to request data from individual sites in January 2017. We received data from 20 different sites across three continents (Europe, North America, and East Asia) and four climate regimes (tropical, temperate, Mediterranean and boreal). We utilized data reported by researchers for climate (annual temperature and precipitation), atmospheric N deposition (wet, bulk, and throughfall), and stream inorganic N concentrations and fluxes across climate regimes and continents, for all years researchers made data available. See Table 1 for a list of publications that include detailed methods utilized by researchers at individual sites. Other site variables, including soil pH, soil C:N ratio, and stand age, may correlate with stream N concentrations and fluxes, but these factors were outside the scope of this study and therefore were not included in our analysis. The ILTER Network is a bottom-up network of networks (Mirtl et al. 2018) where sites decide on their own which environmental variables are measured. However, in recent years large-scale harmonization efforts have begun (Haase et al. 2018). Thirteen of the 20 sites are in Europe and five sites are in the United States (Figure 1). There are two sites from East Asia; both (Japan and Taiwan) are from islands, so no sites are from mainland Asia. Fifteen of the 20 sites are within temperate ecosystems and only two in boreal, one in Mediterranean, and two in tropical ecosystems.

Measurements of wet deposition include the collection and analysis of rainfall and snow, and excludes particulates found in dry deposition. Bulk deposition collectors are similar to those of wet

deposition collectors, but are typically left open, permitting particulates to be collected in the sampler. Throughfall measurements include atmosphere inputs that pass through vegetation canopy, and often include both wet and dry inputs as collectors are typically left open. Study length ranged from 5 to 51 years with data spanning the years 1964 through 2019. We had different sample sizes across variables since not all sites had the same measurements reported.

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All statistical analyses were conducted in R version 4.0.2 (R Core Team 2020). Because of the heterogeneity of our dataset, which includes time-series of different lengths and data-collection methods, we applied a three-step analytical procedure to examine temporal trends. First, we examined the trend of each continuous independent variable (climate and atmospheric N inputs) and response variable (stream N concentrations and fluxes) at each site, using the Mann-Kendall trend test (Kendall 1948, Mann 1945). We used auto- and cross-covariance and correlation functions to identify serially correlated time series (Venerables and Ripley 2002), for which we used the modified Mann-Kendall with the Hamed and Rao (1998) variance correction approach. Serial correlation, also known as temporal autocorrelation, occurs in a time series when a variable is correlated with a lagged version of itself (e.g. a variable at times T and at T+1). This refers to when future observations are affected by past values. This is a common issue in time series, which needs to be considered in analyses like ours (Venerables and Ripley 2002). Second, we tested the relationships between each independent and each response variable at each site (total = 40 combinations per site), using Pearson's correlations. To account for multiple comparisons (at each site, each response variable was tested against six independent variables), we applied a Bonferroni correction to the significance level: alpha = 0.05/6 = 0.0083. We then computed the effect size of the correlations using the function "escal" of the R package "metafor" (Viechtbauer 2010). Third, we ran meta-analysis mixed models (using the R package "metafor", Viechtbauer 2010) and used as response variables the sitespecific S-statistics for the trends of each variable (total = 13 variables) and their variance as effect size of the trends (Kendall 1948, Daufresne et al. 2009, Pilotto et al. 2020), and the effect sizes of the correlations. We ran four types of models: (1) without moderators, to identify the overall patterns of the effect sizes across the whole study area; (2) with "climate regime" (four categories: tropical, temperate, Mediterranean and boreal) as moderator to test how the effect sizes varied among the studied climates; (3) with "continent" (three levels: North America, Europe and Asia) as moderator to test how the effect sizes varied among the studied continents; (4) with "length of the time series" (continuous variable, square-root transformed number of years) as moderator, to test how the effect sizes are affected by the study length. We did not have sufficient replicates across biome types and continents to draw meaningful conclusions about trends of atmospheric deposition or temperature over time for specific climates or continents. For example, there were 15 sites included when examining air temperatures for temperate ecosystems, but only 1-2 boreal, Mediterranean, and tropical sites. In comparing air temperature trends across continents, there were 13 sites in Europe, 5 in North America, and only 2 in Asia. Finally, we ran a sensitivity analysis to evaluate the robustness of the results of the analysis. For that, we identified influential cases (i.e., sites) in each metaanalytical model, following Viechtbauer and Cheung (2010). We then re-ran the models without the influential cases and compared the results with those obtained with the full set of sites. We had neither sufficient sample size nor statistical power to run multivariate analyses to simultaneously determine all primary drivers (e.g., mean annual deposition, temperature, precipitation, continent) and their relative significance for stream N fluxes and concentrations across the 20 sites. We therefore do not refer to climate regime or continent in the remainder of this paper.

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Bulk ammonium, bulk nitrate, and throughfall nitrate decreased across the sites over time (Figure 2; all z-scores < -2; p = 0.038, 0.029, and 0.038, respectively). In contrast, the contribution of ammonium to dissolved inorganic N in wet deposition and throughfall increased significantly over time (Figure 2). The most recent contribution of ammonium to total wet N deposition ranged from 7.3% at H.J. Andrews Forest in Oregon in 2014 to 61.9% in Lago Maggiore area, Italy in 2016. The most recent contribution of ammonium to total throughfall ranged from 36.1% in Krofdorf, Germany in 2013 to 70.2% in Brenna, Poland in 2012. Air temperatures also increased over time with trends most pronounced in Mediterranean and temperate ecosystems and in Europe (Figure 2). Stream nitrate concentrations were positively related to both throughfall ammonium and bulk nitrate deposition fluxes (Figure 3A and C; z-scores >2.2; p = 0.009 and <0.001, respectively). Stream DIN concentrations were positively related to bulk nitrate deposition fluxes (Figure 3D; p =0.001). Rates of annual precipitation were positively related to stream ammonium fluxes (Figure 3E-F; z-score = 3.5; p < 0.001) and stream nitrate fluxes (Figure 4G-H; z-score = 4.0; p < 0.001). Stream nitrate fluxes were positively related to throughfall, wet, and bulk ammonium and nitrate deposition fluxes (Figure 4A-F, H; all z-scores >2.1; p = 0.035, 0.001, 0.022, <0.001, <0.001,and <0.001,respectively). The length of study (5 to 51 years) had both positive and negative effects on the trends we observed (Tables 2 and 3). For example, the longer the study, the stronger the trends we found for increasing temperature over time (Figure 5A) and declining rates of bulk ammonium deposition fluxes over time (Figure 5B). Length of study was positively related to relationships between precipitation and both stream nitrate and DIN concentration (Figure 5E and 5I). There were negative

relationships between length of study and the relationship between temperature and DIN concentration (Figure 5D), precipitation and stream ammonium flux (Figure 5F), and temperature and stream nitrate concentration (Figure 5H). The duration of the study was negatively correlated with the relationship between throughfall nitrate flux and stream ammonium concentration (Figure 5G), and throughfall nitrate flux and stream nitrate concentration (Figure 5J).

Indeed, the declining trends in bulk ammonium and nitrate deposition were most evident in the longest time-series (i.e., negative correlation between these trends and length of study), which has been observed before (Waldner et al. 2014). Similarly, the rapid increases in temperature were best captured with the longest time-series (i.e., positive correlation between trend and length of study).

Our sensitivity analyses showed that excluding influential sites from the analysis did not have statistically significant effects on most of the relationships we report. Influential sites affected the results of only three out of the 106 models (2.8%; 106 models = 13 models to examine trends over time, 13 models to examine potential relationships between trends over time and study length, 40 models to examine correlations, and 40 models to examine potential relationships between correlations and study length) that we ran. Specifically, the Lange Bramke site (a temperate coniferous forest site in Germany) was responsible for the negative relationship between length of study and trend in bulk nitrate deposition, the Brenna site was responsible for the relationship between total inorganic N (DIN) concentration and throughfall ammonium deposition flux, and the TERENO Wüstebach site (another temperate coniferous forest site in Germany; Bogena et al. 2018) drove the relationship between length of study and relationship between nitrate flux and throughfall nitrate deposition flux. In other words, if these sites are excluded from the analysis, these particular relationships are no longer significant.

Patterns Within Sites

Patterns within sites generally mirrored those among all sites (Table S1). For example, air temperature increased significantly at multiple individual sites (i.e., Hubbard Brook Experimental Forest, Krofdorf, Lago Maggiore, Lange Bramke, Parque Natural del Montseny, Piburger See, and River Salaca). Rates of bulk ammonium deposition (Fushan, Lange Bramke, and River Salaca), bulk nitrate deposition (Alptal, Krofdorf, Lange Bramke, and River Salaca), and throughfall nitrate (Krofdorf, Lange Bramke, Zöbelboden, and River Salaca) also declined significantly over time at multiple individual sites (Table S1).

We found statistically significant relationships between atmospheric nitrogen inputs and stream nitrogen concentrations (Table 4) for Brenna (Poland), Hubbard Brook Experimental Forest (U.S.), Krofdorf (Germany), Lago Maggiore (Italy), and Lange Bramke (Germany). We found statistically significant relationships between atmospheric nitrogen inputs and stream nitrogen fluxes for Hubbard Brook Experimental Forest, Krofdorf, Lago Maggiore, Zöbelboden (Austria), Plum Island (U.S.), and TERENO Wüstebach (Germany).

Discussion

Patterns Among Sites

The relationships we found between atmospheric N deposition and stream inorganic N concentrations and fluxes over time and among sites suggest that, despite recent increases in temperature and reductions in total rates of atmospheric N deposition, atmospheric N inputs remain an important driver of stream N cycling in watersheds around the globe. While air temperatures continue to warm, atmospheric deposition and precipitation remain the primary drivers of N export. The observation that rates of bulk and throughfall nitrate deposition declined across the sites over

time was expected given recent government policies that control emissions of N oxides in many, though not all, areas of the globe (Li et al. 2016). In contrast, the pattern of declining bulk ammonium deposition is surprising given that rates remain elevated and even increased in many regions (Templer et al. 2012; Li et al. 2016; EMEP Status Report 2021). Our study (Figure 2) and several recent studies show the trends for declining atmospheric nitrate and the simultaneous increase in contribution of ammonium to N deposition in North America (Li et al. 2016) and Europe (EMEP Status Report 2021). The fact that we did not observe significant patterns in overall wet inorganic N deposition over time (Supplemental Figure 1) was surprising, but this fact is partly attributable to lack of sufficient sample size. Of the 20 datasets included in this study, 11 included measurements of bulk deposition, whereas only eight included wet deposition.

The pattern of increasing temperatures over time across the 20 sites in this data synthesis is consistent with the documented increases in air temperature globally (Hayhoe et al. 2018). In contrast, precipitation is typically more variable than temperature, with precipitation projected to increase in some locations, decrease in others, and generally become more variable throughout the globe (Hayhoe et al. 2018). Our finding that rates of annual precipitation are positively related to stream ammonium and nitrate fluxes align with other studies showing that water inputs can increase N losses through leaching from unmanaged watersheds into nearby streams (Whitehead et al. 2009; Baron et al. 2013). Since N leaching can be strongly driven by runoff events (e.g., snowmelt, stormflows, etc.; Ohte et al. 2004), we might have missed important processes using annual data. As an example, 75% of dissolved inorganic nitrogen is leached during the upper quartile of the discharge in the Zöbelboden catchment, Austria (Dirnböck et al. 2020).

The positive relationships between rates of atmospheric N inputs and stream N concentrations and fluxes across sites support our original hypothesis that sites with greater rates of

atmospheric N deposition have greater N concentrations and fluxes of N in streams. While rates of atmospheric N deposition, especially in the form of nitrate, are declining in many regions around the globe (Figure 2; Lloret and Valiela 2016; Theobald et al. 2019), the magnitude of this decline is small relative to the total amount of N accumulated in the past several decades (Schmitz et al. 2019) and rates of atmospheric N deposition are still increasing or are currently stable in parts of Asia (Ge et al. 2020; Yu et al. 2019). Vuorenmaa et al. (2017) showed that N retention is high in unmanaged forest catchments across Europe even after decades of elevated atmospheric N deposition.

Reforestation or succession after forest harvest should also greatly limit N loss to streams due to inputs of high C:N woody materials (Lajtha 2020, Fisk et al. 2002, Bernhardt et al. 2005, MacDonald et al. 2002). However, a history of elevated N deposition can make forested catchments more prone to an increase in nitrate leaching in the case of partial harvest or other disturbances, as shown for the Alptal site (Schleppi et al., 2017).

Our results also suggest that despite documented reductions in N oxide emissions and nitrate deposition in recent decades, atmospheric N deposition remains a strong regulator of inorganic N exported into nearby aquatic ecosystems. This result is important and timely because it shows that while many government policies are in place to reduce emissions of N oxides and ammonia, many terrestrial ecosystems around the globe are still experiencing the legacy effects of N saturation from past, as well as ongoing, atmospheric N inputs (e.g., Dirnböck et al. 2018) and much of this N is still lost to nearby streams, demonstrating the need for long-term monitoring of both atmospheric N inputs and export to streams. Further, significant reductions in DIN runoff observed in some regions of the globe are likely a result of coupled carbon-N processes (Craine et al. 2018; Groffman et al. 2018). Increasing global temperatures and atmospheric carbon dioxide concentrations, along with

longer growing seasons, lead to increasing demand for N by plants, which may contribute to declining stream N in some ecosystems (Craine et al. 2018; Groffman et al. 2018).

Atmospheric N inputs were related to stream N concentrations and fluxes in the form of nitrate but not ammonium (Figure 3 and 4), which is likely the result of nitrification or ammonium adsorption on clays, or the fact that ammonium is transformed more quickly than nitrate to other forms of N in upland soils, riparian zones, and streams (Peterson et al. 2001; Causse et al. 2015). Although we did not observe any significant changes in stream N concentrations or fluxes over time, the fact that stream nitrate is correlated positively with rates of atmospheric N inputs, even as these inputs are reduced in many regions across the globe, suggests that the legacy effects of past atmospheric N deposition still impact water quality in many locations.

Although the relatively low spatial resolution (i.e., low amount of replication across continents and climates) hindered our ability to draw strong conclusions about climate-specific relationships between atmospheric N inputs and outputs, the relatively long-term record of data we synthesized (data spanned the years 1964 to 2019) from the 20 ILTER sites allowed us to examine trends over time. Our data synthesis also enabled us to examine other variables related to stream N concentrations and fluxes and to determine how these patterns have changed over time. Our results demonstrate the importance of long-term datasets, such as those collected at ILTER sites, to detect significant trends in changes in atmospheric N inputs and stream N over time. Indeed, the declining trends in bulk ammonium and nitrate deposition were most evident in the longest time-series (i.e., negative correlation between these trends and length of study). Similarly, the rapid increases in temperature were best captured with the longest time-series (i.e., positive correlation between trend and length of study). In summary, these results demonstrate that long-term monitoring provides

understanding of ecosystem and biogeochemical cycling over time, that would not be possible with shorter-term studies.

Patterns Within Sites

The consistency of patterns within and across sites further strengthens our findings of increasing air temperatures and declining rates of bulk deposition and throughfall nitrate over time. The strong relationships we observed between atmospheric N inputs and stream nitrogen concentrations and fluxes at individual sites also strengthens our conclusion that atmospheric deposition remains a strong driver of stream nitrogen export in forest and grasslands around the globe.

The location and ecosystem types represented in this data synthesis show that additional datasets pairing atmospheric inputs and stream N output are needed to develop a global understanding of catchment N cycling. Measurements included here are scarce in Central and South America, Africa, Asia, and Oceania, and in arctic, boreal, and tropical biomes, even though ILTER sites are located in most biomes across all continents (Mirtl et al. 2018; Wohner et al. 2021). Our data synthesis included only 20 of the approximately 600 terrestrial research sites across the ILTER network, with the majority (15 out of 20 sites) in temperate ecosystems and only two in boreal, one in Mediterranean, and two in tropical ecosystems. The lack of data from other continents and climates around the globe shows that we have not yet realized the potential for coordinated N research across the ILTER Network. Adopting a harmonized set of environmental monitoring activities with a set of standard variables (including atmospheric deposition samplers) that are recommended to be measured at all ILTER sites (Haase et al. 2018) would enhance the ability to develop global understanding of N dynamics. Furthermore, future data synthesis would be more

effective if these datasets are made publicly available following Findable, Accessible, Interoperable, and Reusable (FAIR) guiding principles (Wilkinson et al. 2016). The observation that longer datasets appear to uncover a larger number of significant relationships than shorter datasets (driven by the fact that environmental variables are inherently variable, masking temporal trends) demonstrates the need for more long-term datasets around the globe.

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Acknowledgements

We thank the organizers and funders of the ILTER Nitrogen Initiative Training Course and Workshop in Hokkaido, Japan in June 2016, which brought together many of the participants in this project. Templer was supported by a US National Science Foundation LTER grant NSF DEB 1637685. McDowell was supported by US National Science Foundation LTER grant NSF DEB 1831592. We are grateful to the EU Horizon 2020 funded eLTER PLUS project (Grand Agreement No. 871128) for financial support to Haase and Dirnböck. Dirnböck was also funded by the LTER-CWN project (FFG project number 858024). This study was partly supported by the Research Initiative Grants of the ILTER, Grants-in-Aid for Scientific Research (17H03833), and Research Institute for Humanity and Nature (RIHN; a constituent member of NIHU) Project No. 14200156. Sharif was supported by NRT-INFEWS: UMD Global STEWARDS (STEM Training at the Nexus of Energy, WAter Reuse and FooD Systems) that was awarded to the University of Maryland School of Public Health by the National Science Foundation National Research Traineeship Program, Grant number 1828910. The monitoring of the Svartberget site in Sweden was funded by the SITES program from the Swedish Research Council. The monitoring of the Volbu Nyhaga site in Norway was part of JOVA - The Norwegian Agricultural Environmental Monitoring Programme, financed by the Ministry of Agriculture and Food. Laitha was supported by NSF grants DEB-1257032 and DEB-

1440409 to the H. J. Andrews Long Term Ecological Research program. The data collection in the Wüstebach catchment was supported by TERENO (Terrestrial Environmental Observatories) funded by the Helmholtz-Gemeinschaft. We thank the "Hessisches Landesamt für Naturschutz, Umwelt und Geologie" for providing data from the Rhine-Main-Observatory. I.Kokorīte was supported by the University of Latvia grant No. AAp2016/B041//Zd2016/AZ03. We thank the Latvian Environment, Geology and Meteorology Center for providing the monitoring data for Latvian site. Anna Avila was supported by Spanish Ministry of Science projects CGL2017-84687-C2-2-R and CGL2009-13188-C03-01. We thank the Tyrolean Alps Long-Term Sociological Ecological and Research (LTSER, Austria) and R. Psenner and S. Morales for helping with data for Piburger See. Data from Hubbard Brook were supported by the National Science Foundation (DEB-1907683) and US Forest Service, Northern Research Station. A. Robison was supported by US National Science Foundation LTER grant NSF OCE 1637630. W. Dodds was supported by NSF DEB 2025849. Observations at Lange Bramke site were funded by the Ministry of Nutrition, Agriculture and Consumer Protection of Lower Saxony under the Permanent Soil Monitoring Programme. The Krofdorf site was funded by the Hessian Ministry of Environment, Climate Protection, Agriculture and Consumer Protection under the "Waldökosystemstudie Hessen". SRM Lins was supported by the São Paulo Research Foundation - FAPESP, grant number 2012/20377-9. Long-term monitoring at Lake Maggiore (LTER site EU-IT08-001-A) was funded by the International Commission for the Protection of Swiss-Italian waters (CIPAIS). We thank Emma Conrad-Rooney for help with Figure 1.

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Declarations

Funding: Please see complete details of funding support in Acknowledgements section above.

451 Conflicts of interest/Competing interests: The authors confirm they do not have any conflict of 452 interest. 453 454 Availability of data and material: We will make our data publicly available through the 455 Environmental Data Initiative (EDI; https://environmentaldatainitiative.org) prior to publication. The 456 EDI Data Repository issues a full dataset citation, including a DOI. 457 458 Code availability: We will make our code publicly available through the Environmental Data 459 Initiative (EDI; https://environmentaldatainitiative.org) prior to publication. The EDI Data 460 Repository issues a full dataset citation, including a DOI. 461 462 Author's contributions: Templer led the data analysis and writing of the manuscript. Harrison and 463 Pilotto analyzed the data. All authors contributed to the writing and editing of the manuscript.

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Figure Legends

798

799 Figure 1. Distribution of sites used in our data synthesis. Triangles represent sites from North America; 800 squares represent sites from Europe; circles represent sites from Asia. 801 Figure 2. A-F. Results of the meta-analysis mixed models are shown for six response variables that 802 changed statistically significantly over time (p < 0.05). G. S-statistic values with 95% confidence 803 intervals shown. S-statistic values that have 95% confidence intervals that do not overlap with zero 804 indicate statistically significant trends over time. Negative values show decreasing trends over time, 805 while positive values show increasing trends over time. N = 10, 11, 8, and 20 sites for bulk 806 ammonium deposition, bulk nitrate deposition, throughfall nitrate, and temperature, respectively. 807 Figure 3. A-E. Relationships between atmospheric N inputs or climate and stream N concentrations or 808 fluxes that were statistically significant (p < 0.05). F. Correlation coefficients shown with 95% 809 confidence intervals. Correlation coefficients that have 95% confidence intervals that do not overlap 810 with zero indicate statistically significant relationships. See Figure 1 for site legend. 811 Figure 4. A-G. Relationships between atmospheric N inputs or climate and stream nitrate flux that were 812 statistically significant (p < 0.05). H. Correlation coefficients shown with 95% confidence intervals. 813 Correlation coefficients that have 95% confidence intervals that do not overlap with zero indicate 814 statistically significant relationships. See Figure 1 for site legend. 815 Figure 5. Association between effect sizes and length of study, as resulting from meta-regressions. The 816 fitted meta-regression lines (blue lines) with 95% confidence intervals (gray shaded area) and the 817 observed values at each study sites (black dots) are shown for significant meta-regressions (p < 818 0.05). Spacing on the x axis follows the square-root scale. A-C: The effect size on the y-axis is S 819 statistics from the Mann-Kendall test. It is a measure of the strength of the monotonic trend of the 820 variable through time at each study site and is positive if the trend is positive and negative if the

trend is negative. D-J: The effect size on the y-axis is the correlation coefficient between the pairs
of variables at each study site.

Supplemental Figure 1. A-B. Wet deposition of ammonium and nitrate over time. Neither of these fluxes
varied over time in a statistically significant way. See Figure 1 for site legend.

Figure 1. Distribution of sites used in our data synthesis. Triangles represent sites from North

America; squares represent sites from Europe; circles represent sites from Asia.



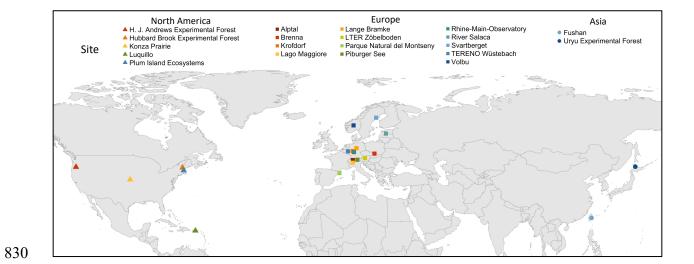


Figure 2. A-F. Results of the meta-analysis mixed models are shown for four response variables that changed statistically significantly over time (P < 0.05). G. S-statistic values with 95% confidence intervals shown. S-statistic values that have 95% confidence intervals that do not overlap with zero indicate statistically significant trends over time. Negative values show decreasing trends over time, while positive values show increasing trends over time. N = 10, 11, 8, 20, 8, and 7 sites for bulk ammonium deposition, bulk nitrate deposition, throughfall nitrate, temperature, and percent ammonium in wet deposition and throughfall, respectively. Percent ammonium in wet deposition and throughfall were calculated as the amount of nitrogen from ammonium relative to DIN (ammonium plus nitrate) at each site. See Figure 1 for site legend.

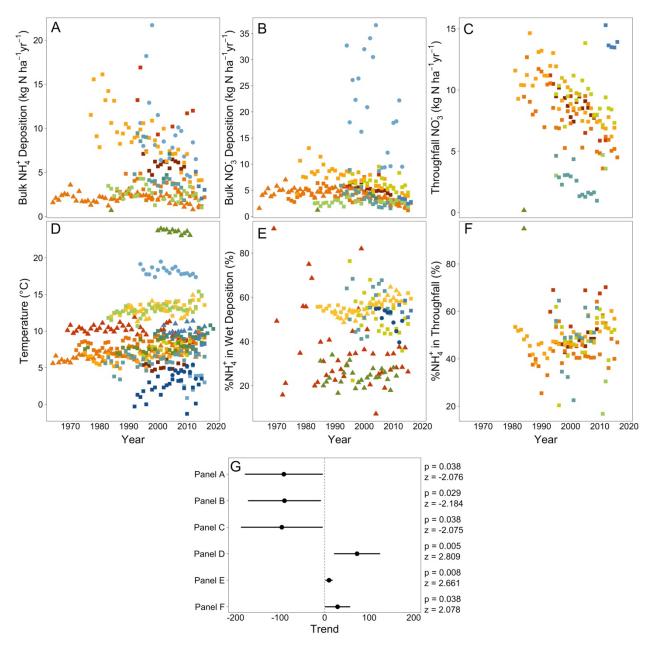


Figure 3. A-E. Relationships between atmospheric N inputs or climate and stream N that were statistically significant (P < 0.05). F. Correlation coefficients shown with 95% confidence intervals. Correlation coefficients that have 95% confidence intervals that do not overlap with zero indicate statistically significant relationships. See Figure 1 for site legend.

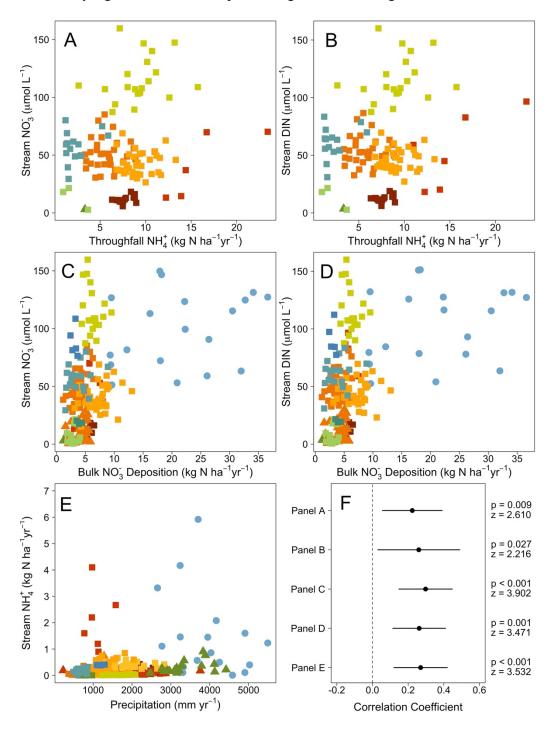


Figure 4. A-G. Relationships between atmospheric N inputs or climate and stream nitrate flux that were statistically significant (P < 0.05). H. Correlation coefficients shown with 95% confidence intervals. Correlation coefficients that have 95% confidence intervals that do not overlap with zero indicate statistically significant relationships. See Figure 1 for site legend.

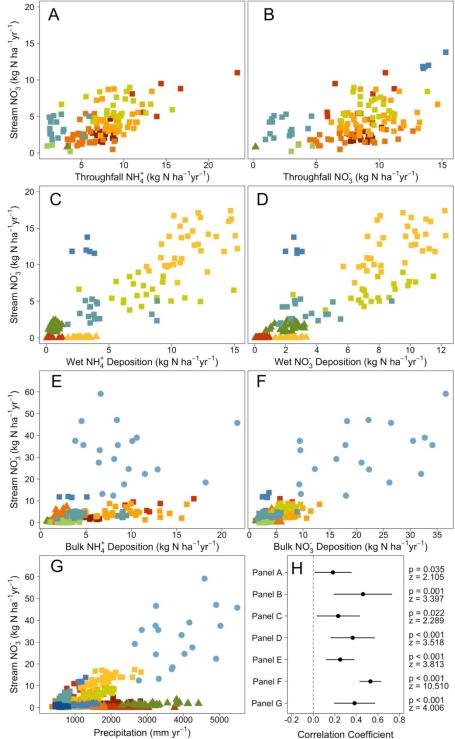
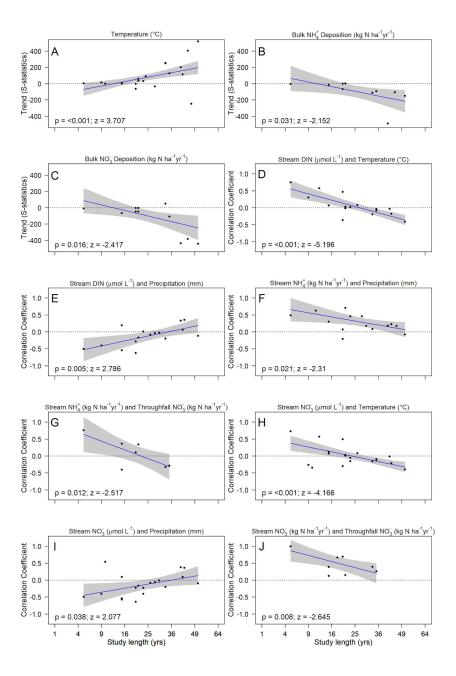


Figure 5. Association between effect sizes and length of study, as resulting from meta-regressions. The fitted meta-regression lines (blue lines) with 95% confidence intervals (gray shaded area) and the observed values at each study sites (black dots) are shown for significant meta-regressions (P < 0.05). Spacing on the x axis follows the square-root scale. A-C: The effect size on the y-axis is S statistics from the Mann-Kendall test. It is a measure of the strength of the monotonic trend of the variable through time at each study site and is positive if the trend is positive and negative if the trend is negative. D-J: The effect size on the y-axis is the correlation coefficient between the pairs of variables at each study site. See Figure 1 for site legend.



| | | | | | | | Mean Annual Temperature | | Start | Study Length | | |
|-----------------------------|----------------|---------------|----------|-----------|---------------|-------------------|----------------------------|------|-------|-----------------|--------------------------------------|--|
| Site | Continent | Country | Latitude | Longitude | Climate | Ecosystem Type | (°) | (mm) | Year | (yr) | DEIMS | Citation |
| Alptal | Europe | Switzerland | 47.0440 | 8.7130 | Temperate | Coniferous Forest | 5.3 | 2243 | 1995 | 15 | 9e1c8ec8-a407-426a-8410-05180b96e75a | Schleppi et al 2017 |
| Brenna | Europe | Poland | 49.6604 | 18.9371 | Temperate | Coniferous Forest | 8.7 | 1091 | 1993 | 20 | 0ff5485d-4436-4153-b6fb-d6eac9c9dd23 | |
| Fushan | Asia* | Taiwan | 23.5667 | 121.5667 | Tropical | Evergreen Forest | 18.2 | 3862 | 1994 | 20 | 14c64e04-4152-455h-97a2-2cec1d2af13c | Chang et al 2017 |
| HJ Andrews Experimental For | North America | United States | 44.2323 | -122.1690 | Temperate | Coniferous Forest | 10.0 | 2210 | 1969 | 47 | a7136f22-6d82-4177-b85a-82713b6fff5e | Vanderbilt et al. 2003; Lajtha & Jones 2018 |
| Hubbard Brook Exper For | North America | United States | 43.9620 | -71.8050 | Temperate | Broadleaf Forest | 6.7 | 1448 | 1964 | 51 | 635846c4-e015-431a-8211-009e7785b4b6 | HBWatER 2021a & 2021b |
| Konza Prairie | North America | United States | 39.0954 | -96.5750 | Temperate | Grassland | 12.8 | 849 | 1990 | 23 | a635d4dc-6a74-4947-b3d1-5c88b3a441c7 | Dodds 2020; Kemp & Doods 2001 |
| Krofdorf | Europe | Germany | 50.6800 | 8.6500 | Temperate | Broadleaf Forest | 8.5 | 705 | 1972 | 45 | f73a0f95-8fb0-4755-92fc-f4b0207f5fe4 | Schumann et al 2010 |
| Lago Maggiore | Europe | Italy | 45.9547 | 8.6340 | Temperate | Mixed Forest | 13.2 | 1827 | 1988 | 33 | f30007c4-8a6e-4f11-ab87-569db54638fe | Rogora et al 2007 |
| Lange Bramke | Europe | Germany | 51.8667 | 10.4333 | Temperate | Coniferous Forest | 6.5 | 1308 | 1975 | 41 | 8e24d4f8-d6f6-4463-83e9-73cac2fd3f38 | Meesenburg et al 2010; Ahrends et al 2010 |
| LTER Zöbelboden | Europe | Austria | 47.8422 | 14.4441 | Temperate | Mixed Forest | 7.9 | 1649 | 1995 | 21 | 8eda49e9-1f4e-4f3e-b58e-e0bb25dc32a6 | Dirnböck et al 2020 |
| Luquillo | North America* | United States | 18.3607 | -65.7015 | Tropical | Rainforest | 23.6 | 3204 | 1984 | 28 | bd0b5bcf-4f2e-4038-8275-629ffa5bf2aa | McDowell et al 2021 |
| Parque Natural del Montseny | Europe | Spain | 41.7773 | -2.3528 | Mediterranean | Forest/Woodland | 13.0 | 866 | 1983 | 33 | 19fd543e-53b2-478e-b9c8-7d1160a0ee82 | Ávila et al 2020 |
| Piburger See | Europe | Austria | 47.1833 | 10.8833 | Temperate | Coniferous Forest | 7.8 | 750 | 1975 | 42 | ed1f621f-d337-4a3c-9cf1-7be144fc556c | Niedrist et al 2018 |
| Plum Island Ecosystems | North America | United States | 42.8276 | -71.2198 | Temperate | Mixed Forest | 10.2 | 1277 | 2002 | 15 | 4c2cfbde-0fcf-4a87-99ee-52182bfa240a | Morse and Wollheim 2014 |
| Rhine-Main-Observatory | Europe | Germany | 50.2673 | 9.2691 | Temperate | Grassland | 9.5 | 813 | 1999 | 21 | 9f9ba137-342d-4813-ae58-a60911c3abc1 | Kuemmerlen et al 2015 |
| River Salaca | Europe | Latvia | 57.8000 | 24.3333 | Temperate | Coniferous Forest | 6.0 | 690 | 1982 | 35 | 81a2b50d-76ea-426b-8e5b-0560dc07ee57 | Terauda & Nikodemus 2007; Melecis et al 2005 |
| Svartberget | Europe | Sweden | 64.2372 | 19.7611 | Boreal | Boreal forest | 3.0 | 640 | 2008 | 9 | c0705d0f-92c1-4964-a345-38c0be3113e1 | Laudon et al 2013 |
| TERENO Wüstebach | Europe | Germany | 50.5833 | 6.4333 | Temperate | Coniferous Forest | 8.1 | 1170 | 2012 | 5 | 9fe5a5d1-ccc0-41ab-b555-5ca44da24cd8 | Bogena et al 2018 |
| Uryu Experimental For | Asia* | Japan | 44.3650 | 142.2610 | Temperate | Mixed Forest | 4.5 | 1435 | 2004 | 10 | e7bc527e-1dd1-4898-b373-e3cd3b4596bb | Fukuzawa et al 2020; NARD |
| Volbu Nyhaga | Europe | Norway | 61.1200 | 9.0600 | Boreal | Boreal forest | 2.2 | 607 | 1992 | 24 | a7e1e6e2-6275-4cb6-855e-f2f7aa79cfa1 | Hauken et al 2020 |

Table 1. Sites included in our data synthesis. "Mixed Forest" includes broadleaf and conifer trees. The study length refers to the maximum length of time for response variables within a site; note that the study length may actually be shorter for some variables within a site. Dynamic Ecological Information Management System (DEIMS; deims.org) numbers included for all sites. *Fushan, Luquillo, and Uryu Experimental Forest are located on islands and not mainland

continents.

Table 2. Results of the meta-regressions testing the general patterns in the temporal trends of the response variables and of the explanatory variables. Response and explanatory variables are to be intended as temporal trends (Mann-Kendall S-statistics); z = test statistics of the coefficient, p = p-values for the test statistics. Percent ammonium in wet, throughfall, and bulk deposition were calculated as the amount of nitrogen from ammonium relative to DIN (ammonium plus nitrate) at each site in all three forms of atmospheric N inputs. Statistically significant results are highlighted in bold. The "Overall" analysis shows if there is an overall pattern in the trends (positive z: increasing trends, negative z: decreasing trends) across all study sites.

| | | Ove | erall | Length of | Time Series |
|---|----|--------|-------|-----------|-------------|
| | n | z | p | Z | p |
| Stream NH ₄ (mmol L ⁻¹) | 16 | -0.997 | 0.319 | -1.399 | 0.162 |
| Stream NO ₃ (mmol L ⁻¹) | 19 | -0.479 | 0.632 | -0.365 | 0.715 |
| Stream NH ₄ ⁺ (kg N ha ⁻¹ yr ⁻¹) | 14 | -0.333 | 0.739 | -0.459 | 0.646 |
| Stream NO ₃ - (kg N ha ⁻¹ yr ⁻¹) | 17 | -1.507 | 0.132 | -0.942 | 0.346 |
| Stream DIN (mmol L ⁻¹) | 16 | -0.197 | 0.844 | -0.374 | 0.709 |
| Precipitation (mm) | 20 | 0.799 | 0.424 | 0.921 | 0.357 |
| Temperature (°C) | 20 | 2.809 | 0.005 | 3.707 | 0.001 |
| Wet NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 8 | -0.190 | 0.849 | -0.402 | 0.688 |
| Wet NO ₃ Deposition (kg N ha ⁻¹ yr ⁻¹) | 8 | 0.033 | 0.973 | 0.214 | 0.831 |
| Wet %NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 8 | 2.661 | 0.008 | 0.042 | 0.967 |
| Throughfall NH ₄ ⁺ (kg N ha ⁻¹ yr ⁻¹) | 7 | -0.994 | 0.320 | -1.439 | 0.150 |
| Throughfall NO ₃ - (kg N ha ⁻¹ yr ⁻¹) | 8 | -2.075 | 0.038 | -1.927 | 0.054 |
| Throughfall %NH ₄ (kg N ha ⁻¹ yr ⁻¹) | 7 | 2.078 | 0.038 | 0.912 | 0.362 |
| Bulk NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 10 | -2.076 | 0.038 | -2.152 | 0.031 |
| Bulk NO ₃ Deposition (kg N ha-1 yr-1) | 11 | -2.184 | 0.029 | -2.417 | 0.016 |
| Bulk %NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 10 | -0.022 | 0.982 | 0.299 | 0.765 |

Table 3. Results of the meta-regressions testing the general patterns in the correlations between pairs of response and explanatory variables. z = test statistics of the coefficient, p = p-values for the test statistics. Statistically significant results are highlighted in bold. The "Overall" analysis shows if there is an overall pattern in the correlations between the two variables (positive z: positive correlation, negative z: negative correlation) across all study sites.

| | | | Overall | | Length of | Time Series |
|--|--|----|---------|-------|-----------|-------------|
| Variable 1 | Variable 2 | n | z | p | z | p |
| Stream NH ₄ ⁺ (µmol L ⁻¹) | Temperature (°C) | 15 | 0.254 | 0.800 | -1.695 | 0.090 |
| | Bulk NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 9 | 0.722 | 0.470 | -0.800 | 0.424 |
| | Throughfall NH ₄ (kg N ha ⁻¹ yr ⁻¹) | 6 | 0.240 | 0.811 | -0.305 | 0.761 |
| | Wet NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 6 | 0.250 | 0.802 | -0.300 | 0.764 |
| | Bulk NO ₃ - Deposition (kg N ha ⁻¹ yr ⁻¹) | 10 | -1.305 | 0.192 | 0.183 | 0.855 |
| | Throughfall NO ₃ (kg N ha ⁻¹ yr ⁻¹) | 7 | -0.625 | 0.532 | 0.216 | 0.829 |
| | Wet NO ₃ - Deposition (kg N ha ⁻¹ yr ⁻¹) | 6 | -1.452 | 0.146 | -0.160 | 0.873 |
| | Precipitation (mm) | 15 | -1.098 | 0.272 | -1.287 | 0.198 |
| Stream NH ₄ (kg N ha ⁻¹ yr ⁻¹) | Temperature (°C) | 14 | -0.393 | 0.694 | -0.230 | 0.818 |
| | Bulk NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 8 | 1.138 | 0.255 | -0.685 | 0.493 |
| | Throughfall NH ₄ (kg N ha ⁻¹ yr ⁻¹) | 6 | -0.136 | 0.892 | -0.275 | 0.783 |
| | Wet NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 7 | 1.488 | 0.137 | -1.516 | 0.129 |
| | Bulk NO ₃ Deposition (kg N ha ⁻¹ yr ⁻¹) | 9 | 0.843 | 0.399 | -1.653 | 0.098 |
| | Throughfall NO ₃ (kg N ha ⁻¹ yr ⁻¹) | 7 | 0.376 | 0.707 | -2.517 | 0.012 |
| | Wet NO ₃ Deposition (kg N ha ⁻¹ yr ⁻¹) | 7 | 1.300 | 0.194 | -0.201 | 0.841 |
| | Precipitation (mm) | 14 | 3.532 | 0.000 | -2.310 | 0.021 |
| Stream DIN (µmol L ⁻¹) | Temperature (°C) | 16 | 0.265 | 0.791 | -5.196 | 0.000 |
| | Bulk NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 10 | 1.190 | 0.234 | -0.527 | 0.598 |
| | Throughfall NH ₄ (kg N ha ⁻¹ yr ⁻¹) | 7 | 2.216 | 0.027 | -1.793 | 0.073 |
| | Wet NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 6 | 0.227 | 0.821 | -0.542 | 0.588 |
| | Bulk NO ₃ Deposition (kg N ha ⁻¹ yr ⁻¹) | 11 | 3.471 | 0.001 | 1.656 | 0.098 |
| | Throughfall NO ₃ (kg N ha ⁻¹ yr ⁻¹) | 8 | 0.870 | 0.385 | 0.987 | 0.324 |
| | Wet NO ₃ Deposition (kg N ha ⁻¹ yr ⁻¹) | 6 | -0.141 | 0.888 | -0.283 | 0.777 |
| | Precipitation (mm) | 16 | -1.135 | 0.256 | 2.786 | 0.005 |
| Stream NO ₃ - (µmol L-1) | Temperature (°C) | 19 | -0.379 | 0.704 | -4.166 | 0.000 |
| | Bulk NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 10 | 1.157 | 0.247 | -0.228 | 0.820 |
| | Throughfall NH ₄ (kg N ha ⁻¹ yr ⁻¹) | 7 | 2.610 | 0.009 | -1.657 | 0.097 |
| | Wet NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 7 | -0.054 | 0.957 | -0.600 | 0.548 |
| | Bulk NO ₃ Deposition (kg N ha ⁻¹ yr ⁻¹) | 11 | 3.902 | 0.000 | 1.707 | 0.088 |
| | Throughfall NO ₃ (kg N ha ⁻¹ yr ⁻¹) | 8 | 1.074 | 0.283 | 1.085 | 0.278 |
| | Wet NO ₃ Deposition (kg N ha ⁻¹ yr ⁻¹) | 7 | 0.650 | 0.516 | -0.864 | 0.388 |
| | Precipitation (mm) | 19 | -1.945 | 0.052 | 2.077 | 0.038 |
| Stream NO ₃ (kg N ha ⁻¹ yr ⁻¹) | Temperature (°C) | 17 | -0.770 | 0.441 | -0.449 | 0.653 |
| | Bulk NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 9 | 3.813 | 0.000 | 0.032 | 0.975 |
| | Throughfall NH ₄ ⁺ (kg N ha ⁻¹ yr ⁻¹) | 7 | 2.105 | 0.035 | -0.417 | 0.677 |
| | Wet NH ₄ ⁺ Deposition (kg N ha ⁻¹ yr ⁻¹) | 7 | 2.289 | 0.022 | 0.290 | 0.772 |
| | Bulk NO ₃ Deposition (kg N ha ⁻¹ yr ⁻¹) | 10 | 10.510 | 0.000 | 0.067 | 0.947 |
| | Throughfall NO ₃ (kg N ha ⁻¹ yr ⁻¹) | 8 | 3.397 | 0.001 | -2.645 | 0.008 |
| | Wet NO ₃ Deposition (kg N ha ⁻¹ yr ⁻¹) | 7 | 3.518 | 0.000 | -0.162 | 0.871 |
| | Precipitation (mm) | 17 | 4.006 | 0.000 | -1.463 | 0.143 |

Table 4. Significance of the correlation between response and independent variables at each site. Only those sites with statistically significant relationships are shown. To account for multiple comparisons (at each site, each response variable was tested against six independent variables), we applied a Bonferroni correction to the significance level: alpha = 0.05/6 = 0.0083.

| Site | Deposition or Climate Variable | Stream Variable | n | r | р |
|-------------------------|--|-----------------------------------|-----|--------|-------|
| Brenna | Throughfall NH ₄ ⁺ | NH ₄ Concentration | 6 | 0.973 | 0.001 |
| Brenna | Throughfall NO ₃ | NO ₃ Concentration | 6 | 0.965 | 0.002 |
| Brenna | Throughfall NO ₃ | DIN Concentration | 6 | 0.939 | 0.005 |
| Hubbard Brook Exper For | Temperature | NO ₃ Concentration | 51 | -0.395 | 0.004 |
| Hubbard Brook Exper For | Bulk NO ₃ | NO ₃ Concentration | 51 | 0.406 | 0.003 |
| Hubbard Brook Exper For | Bulk NO ₃ | NO ₃ - Flux | 51 | 0.473 | 0.000 |
| Hubbard Brook Exper For | Temperature | DIN Concentration | 51 | -0.411 | 0.003 |
| Hubbard Brook Exper For | Bulk NO ₃ | DIN Concentration | 51 | 0.404 | 0.003 |
| Krofdorf | Bulk NO ₃ | NH ₄ Concentration | 43 | -0.462 | 0.002 |
| Krofdorf | Bulk NO ₃ | NO ₃ Concentration | 43 | 0.530 | 0.000 |
| Krofdorf | Bulk NO ₃ | NO ₃ Flux | 43 | 0.666 | 0.000 |
| Krofdorf | Bulk NO ₃ | DIN Concentration | 43 | 0.470 | 0.001 |
| Lago Maggiore | Precipitation | NH ₄ Concentration | 121 | -0.406 | 0.000 |
| Lago Maggiore | Temperature | NH ₄ Concentration | 123 | -0.443 | 0.000 |
| Lago Maggiore | Wet NH ₄ ⁺ | NH ₄ Concentration | 123 | -0.431 | 0.000 |
| Lago Maggiore | Wet NO ₃ | NH ₄ Concentration | 123 | -0.368 | 0.000 |
| Lago Maggiore | Temperature | NO ₃ Concentration | 126 | 0.261 | 0.003 |
| Lago Maggiore | Wet NH ₄ ⁺ | NH ₄ Concentration | 126 | 0.557 | 0.000 |
| Lago Maggiore | Wet NO ₃ | NO ₃ Concentration | 126 | 0.500 | 0.000 |
| Lago Maggiore | Temperature | NH ₄ Flux | 123 | -0.399 | 0.000 |
| Lago Maggiore | Wet NH ₄ ⁺ | NH ₄ Flux | 123 | -0.419 | 0.000 |
| Lago Maggiore | Wet NO ₃ | NH ₄ Flux | 123 | -0.354 | 0.000 |
| Lago Maggiore | Precipitation | NO ₃ Flux | 124 | 0.437 | 0.000 |
| Lago Maggiore | Wet NH ₄ ⁺ | NO ₃ Flux | 126 | 0.588 | 0.000 |
| Lago Maggiore | Wet NO ₃ | NO ₃ Flux | 126 | 0.620 | 0.000 |
| Lago Maggiore | Wet NH ₄ ⁺ | DIN Concentration | 123 | 0.524 | 0.000 |
| Lago Maggiore | Wet NO ₃ | DIN Concentration | 123 | 0.471 | 0.000 |
| Lange Bramke | Throughfall NO ₃ | NH ₄ Concentration | 35 | -0.462 | 0.005 |
| Lange Bramke | Bulk NO ₃ | NH ₄ Concentration | 39 | -0.447 | 0.004 |
| Lange Bramke | Precipitation | NO ₃ Flux | 41 | 0.755 | 0.000 |
| LTER Zöbelboden | Precipitation | NH ₄ ⁺ Flux | 21 | 0.709 | 0.000 |
| LTER Zöbelboden | Precipitation | NO ₃ Flux | 21 | 0.693 | 0.000 |
| LTER Zöbelboden | Wet NO ₃ | NO ₃ - Flux | 21 | 0.644 | 0.002 |
| Plum Island Ecosystems | Precipitation | NO ₃ Flux | 30 | 0.774 | 0.000 |
| TERENO Wüstebach | Throughfall NO ₃ | NO ₃ Flux | 5 | 0.988 | 0.001 |

Table S1. Results of the Mann-Kendall trend test at each site, which shows which response variable changed significantly over time. The Hamed and Rao (1998) variance correction approach was applied to serially correlated time series. Relationships that are statistically significant (p < 0.0083) are in bold. We applied a Bonferroni correction to the significance level: alpha = 0.05/6 = 0.0083.

| | Stream 1 | NH4+ Conc | entration | Stream | NO3- Conc | entration | Str | eam NH4+ | Flux | Str | eam NO3- I | Flux |
|---|----------|-----------|-----------|--------|-----------|-----------|--------|----------|---------|--------|------------|---------|
| Site | S | Var | P-Value | S | Var | P-Value | S | Var | P-Value | S | Var | P-Value |
| Alptal | -7.0 | 408.3 | 0.77 | -39.0 | 408.3 | 0.06 | -15.0 | 253.7 | 0.38 | -60.0 | 407.3 | 0.00 |
| Brenna | -9.0 | 28.3 | 0.13 | -11.0 | 28.3 | 0.06 | 1.0 | 28.3 | 1.00 | -9.0 | 28.3 | 0.13 |
| Fushan | -20.0 | 1384.1 | 0.61 | 33.0 | 817.0 | 0.26 | -28.0 | 813.3 | 0.34 | 37.0 | 817.0 | 0.21 |
| H. J. Andrews Experimental Forest | NA | NA | NA | NA | NA | NA | 75.0 | 3232.1 | 0.19 | -139.0 | 11268.1 | 0.19 |
| Hubbard Brook Exper For | -732.0 | 6712.7 | 0.00 | -713.0 | 41384.3 | 0.00 | -677.0 | 11157.1 | 0.00 | -641.0 | 40990.4 | 0.00 |
| Konza Prairie | 155.0 | 1098.5 | 0.00 | -3.0 | 1433.7 | 0.96 | 127.0 | 1425.0 | 0.00 | -27.0 | 1433.7 | 0.49 |
| Krofdorf | 333.0 | 12261.5 | 0.00 | -93.0 | 17775.4 | 0.49 | 362.0 | 8915.3 | 0.00 | -194.0 | 15074.2 | 0.12 |
| Lago Maggiore | -227.0 | 6783.5 | 0.01 | -92.0 | 5281.3 | 0.21 | -141.0 | 4162.3 | 0.03 | -132.0 | 1364.1 | 0.00 |
| Lange Bramke | 228.0 | 18652.1 | 0.10 | 155.0 | 18540.5 | 0.26 | 188.0 | 18775.7 | 0.17 | 158.0 | 7926.7 | 0.08 |
| LTER Zöbelboden | -15.0 | 1978.7 | 0.75 | 74.0 | 2034.2 | 0.11 | -10.0 | 1090.0 | 0.79 | 48.0 | 1096.7 | 0.16 |
| Luquillo | -41.0 | 817.0 | 0.16 | -136.0 | 6663.8 | 0.10 | -11.0 | 528.0 | 0.66 | -14.0 | 2562.0 | 0.80 |
| Parque Natural del Montseny | 0.0 | 0.0 | NA | 39.0 | 408.3 | 0.06 | NA | NA | NA | 24.0 | 1254.7 | 0.52 |
| Piburger See | -309.0 | 4157.7 | 0.00 | 310.0 | 3843.3 | 0.00 | 28.0 | 162.0 | 0.03 | 27.0 | 165.0 | 0.04 |
| Plum Island Ecosystems | NA | NA | NA | 0.0 | 145.3 | 0.93 | NA | NA | NA | -37.0 | 408.3 | 0.07 |
| Rhine-Main-Observatory | -25.0 | 1095.7 | 0.47 | -107.0 | 1095.7 | 0.00 | NA | NA | NA | NA | NA | NA |
| River Salaca | -216.0 | 11756.1 | 0.05 | -241.0 | 6388.9 | 0.00 | -195.0 | 10862.5 | 0.06 | -195.0 | 4958.3 | 0.01 |
| Svartberget field infrastructure (LTER) | 2.0 | 92.0 | 0.92 | -6.0 | 92.0 | 0.60 | NA | NA | NA | NA | NA | NA |
| TERENO Wüstebach | 6.0 | 16.7 | 0.22 | 2.0 | 16.7 | 0.81 | -3.0 | 15.7 | 0.61 | 0.0 | 16.7 | 0.81 |
| Uryu Experimental For | NA | NA | NA | 19.0 | 125.0 | 0.11 | NA | NA | NA | NA | NA | NA |
| Volbu | NA | NA | NA | -11.0 | 721.7 | 0.71 | NA | NA | NA | -32.0 | 1432.7 | 0.41 |

| | Stream | DIN Conce | entration | | Precipitation | n | | Temperatur | e | Wet | Deposition | NH4+ |
|---|--------|-----------|-----------|--------|---------------|---------|--------|------------|---------|-------|------------|---------|
| Site | S | Var | P-Value | S | Var | P-Value | S | Var | P-Value | S | Var | P-Value |
| Alptal | -44.0 | 407.3 | 0.03 | -1.0 | 408.3 | 1.00 | 3.0 | 112.6 | 0.85 | NA | NA | NA |
| Brenna | -11.0 | 28.3 | 0.06 | 5.0 | 28.3 | 0.45 | -2.0 | 27.3 | 0.85 | NA | NA | NA |
| Fushan | 31.0 | 817.0 | 0.29 | 8.0 | 950.0 | 0.82 | -64.0 | 950.0 | 0.04 | NA | NA | NA |
| H. J. Andrews Experimental Forest | NA | NA | NA | -106.0 | 1579.3 | 0.01 | -248.0 | 23436.6 | 0.11 | -7.0 | 8487.0 | 0.95 |
| Hubbard Brook Exper For | -739.0 | 38967.0 | 0.00 | 243.0 | 7581.0 | 0.01 | 519.0 | 15158.3 | 0.00 | NA | NA | NA |
| Konza Prairie | 111.0 | 1433.7 | 0.00 | 13.0 | 1433.7 | 0.75 | 32.0 | 1426.0 | 0.41 | 82.0 | 1432.7 | 0.03 |
| Krofdorf | -53.0 | 22020.8 | 0.73 | 128.0 | 10445.3 | 0.21 | 405.0 | 13464.1 | 0.00 | NA | NA | NA |
| Lago Maggiore | -92.0 | 5354.3 | 0.21 | 62.0 | 1737.1 | 0.14 | 254.0 | 4165.3 | 0.00 | -81.0 | 4164.3 | 0.22 |
| Lange Bramke | 200.0 | 18506.2 | 0.14 | -90.0 | 7926.7 | 0.32 | 200.0 | 2645.7 | 0.00 | NA | NA | NA |
| LTER Zöbelboden | 75.0 | 2066.5 | 0.10 | 8.0 | 1096.7 | 0.83 | 60.0 | 1096.7 | 0.07 | -38.0 | 1096.7 | 0.26 |
| Luquillo | -69.0 | 817.0 | 0.02 | 14.0 | 2560.0 | 0.80 | -34.0 | 212.7 | 0.02 | 89.0 | 2054.3 | 0.05 |
| Parque Natural del Montseny | 39.0 | 408.3 | 0.06 | 71.0 | 6578.9 | 0.39 | 249.0 | 2526.9 | 0.00 | NA | NA | NA |
| Piburger See | 300.0 | 4146.0 | 0.00 | 112.0 | 3802.7 | 0.07 | 114.0 | 794.1 | 0.00 | NA | NA | NA |
| Plum Island Ecosystems | NA | NA | NA | -33.0 | 408.3 | 0.11 | 9.0 | 408.3 | 0.69 | NA | NA | NA |
| Rhine-Main-Observatory | -94.0 | 1096.7 | 0.00 | -11.0 | 408.3 | 0.62 | 34.0 | 811.8 | 0.25 | NA | NA | NA |
| River Salaca | -253.0 | 8669.3 | 0.01 | 61.0 | 3141.7 | 0.28 | 127.0 | 1165.4 | 0.00 | -66.0 | 388.6 | 0.00 |
| Svartberget field infrastructure (LTER) | 0.0 | 92.0 | 0.92 | -10.0 | 92.0 | 0.35 | 14.0 | 92.0 | 0.18 | NA | NA | NA |
| TERENO Wüstebach | 4.0 | 16.7 | 0.46 | 0.0 | 16.7 | 0.81 | 4.0 | 16.7 | 0.46 | -2.0 | 16.7 | 0.81 |
| Uryu Experimental For | NA | NA | NA | 27.0 | 125.0 | 0.02 | -1.0 | 15.6 | 1.00 | -1.0 | 125.0 | 1.00 |
| Volbu | NA | NA | NA | 82.0 | 1625.3 | 0.04 | 90.0 | 1620.7 | 0.03 | NA | NA | NA |

| | Wet | Deposition : | NO3- | Dry | Deposition | NH4+ | Dry 1 | Deposition | NO3- | Thr | oughfall N | H4+ |
|---|--------|--------------|---------|-----|------------|---------|-------|------------|---------|--------|------------|---------|
| Site | S | Var | P-Value | S | Var | P-Value | S | Var | P-Value | S | Var | P-Value |
| Alptal | NA | NA | NA | NA | NA | NA | NA | NA | NA | 4.0 | 297.4 | 0.86 |
| Brenna | NA | NA | NA | NA | NA | NA | NA | NA | NA | -5.0 | 28.3 | 0.45 |
| Fushan | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| H. J. Andrews Experimental Forest | 257.0 | 10173.9 | 0.01 | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Hubbard Brook Exper For | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Konza Prairie | -79.0 | 1433.7 | 0.04 | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Krofdorf | NA | NA | NA | NA | NA | NA | NA | NA | NA | -87.0 | 3461.7 | 0.14 |
| Lago Maggiore | -174.0 | 4165.3 | 0.01 | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Lange Bramke | NA | NA | NA | NA | NA | NA | NA | NA | NA | -164.0 | 4088.5 | 0.01 |
| LTER Zöbelboden | 30.0 | 1096.7 | 0.38 | NA | NA | NA | NA | NA | NA | 10.0 | 1096.7 | 0.79 |
| Luquillo | 131.0 | 2058.3 | 0.00 | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Parque Natural del Montseny | NA | NA | NA | NA | NA | NA | NA | NA | NA | -1.0 | 3.7 | 1.00 |
| Piburger See | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Plum Island Ecosystems | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Rhine-Main-Observatory | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| River Salaca | -88.0 | 493.3 | 0.00 | 5.0 | 489.7 | 0.86 | -69.0 | 492.3 | 0.00 | -27.0 | 408.3 | 0.20 |
| Svartberget field infrastructure (LTER) | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| TERENO Wüstebach | -1.0 | 15.7 | 1.00 | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Uryu Experimental For | 25.0 | 125.0 | 0.03 | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Volbu | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |

| | Thi | oughfall N | Ю3- | Bulk | Deposition | NH4+ | Bulk | Deposition | NO3- | Bulk I | Deposition 9 | %NH4+ |
|---|--------|------------|---------|--------|------------|---------|--------|------------|---------|--------|--------------|---------|
| Site | S | Var | P-Value | S | Var | P-Value | S | Var | P-Value | S | Var | P-Value |
| Alptal | -49.0 | 408.3 | 0.02 | -14.0 | 407.3 | 0.52 | -65.0 | 408.3 | 0.00 | 49.0 | 408.3 | 0.02 |
| Brenna | -13.0 | 28.3 | 0.02 | -1.0 | 28.3 | 1.00 | -2.0 | 27.3 | 0.85 | 5.0 | 28.3 | 0.45 |
| Fushan | NA | NA | NA | -68.0 | 414.4 | 0.00 | -46.0 | 539.8 | 0.05 | -2.0 | 898.3 | 0.97 |
| H. J. Andrews Experimental Forest | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Hubbard Brook Exper For | NA | NA | NA | -151.0 | 27931.2 | 0.37 | -443.0 | 60495.7 | 0.07 | 379.0 | 23443.0 | 0.01 |
| Konza Prairie | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Krofdorf | -279.0 | 4164.3 | 0.00 | -107.0 | 3459.7 | 0.07 | -384.0 | 13988.2 | 0.00 | 129.0 | 2296.2 | 0.01 |
| Lago Maggiore | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Lange Bramke | -386.0 | 3762.5 | 0.00 | -489.0 | 4263.9 | 0.00 | -436.0 | 6832.7 | 0.00 | -147.0 | 5096.8 | 0.04 |
| LTER Zöbelboden | -73.0 | 1095.7 | 0.03 | NA | NA | NA | -46.0 | 1096.7 | 0.17 | NA | NA | NA |
| Luquillo | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Parque Natural del Montseny | 1.0 | 3.7 | 1.00 | -113.0 | 3983.0 | 0.08 | 50.0 | 2842.0 | 0.36 | -192.0 | 284.0 | 0.00 |
| Piburger See | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Plum Island Ecosystems | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Rhine-Main-Observatory | NA | NA | NA | 3.0 | 3.7 | 0.30 | -3.0 | 3.7 | 0.30 | 3.0 | 3.7 | 0.30 |
| River Salaca | -55.0 | 408.3 | 0.01 | -93.0 | 697.0 | 0.00 | -109.0 | 697.0 | 0.00 | 17.0 | 697.0 | 0.54 |
| Svartberget field infrastructure (LTER) | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| TERENO Wüstebach | -4.0 | 16.7 | 0.46 | -4.0 | 8.7 | 0.31 | -8.0 | 16.7 | 0.09 | -2.0 | 8.7 | 0.73 |
| Uryu Experimental For | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |
| Volbu | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA | NA |

| | Thro | oughfall %l | NH4+ | Wet D | eposition % | 6NH4+ |
|---|-------|-------------|---------|-------|-------------|---------|
| Site | S | Var | P-Value | S | Var | P-Value |
| Alptal | 45.0 | 408.3 | 0.03 | NA | NA | NA |
| Brenna | 5.0 | 4.2 | 0.05 | NA | NA | NA |
| Fushan | NA | NA | NA | NA | NA | NA |
| H. J. Andrews Experimental Forest | NA | NA | NA | -85.0 | 8514.3 | 0.36 |
| Hubbard Brook Exper For | NA | NA | NA | NA | NA | NA |
| Konza Prairie | NA | NA | NA | 165.0 | 1433.7 | 0.00 |
| Krofdorf | 95.0 | 1759.2 | 0.03 | NA | NA | NA |
| Lago Maggiore | NA | NA | NA | 226.0 | 4165.3 | 0.00 |
| Lange Bramke | 217.0 | 4958.3 | 0.00 | NA | NA | NA |
| LTER Zöbelboden | 40.0 | 1096.7 | 0.24 | -70.0 | 1096.7 | 0.04 |
| Luquillo | NA | NA | NA | 18.0 | 2057.3 | 0.71 |
| Parque Natural del Montseny | -1.0 | 3.7 | 1.00 | NA | NA | NA |
| Piburger See | NA | NA | NA | NA | NA | NA |
| Plum Island Ecosystems | NA | NA | NA | NA | NA | NA |
| Rhine-Main-Observatory | NA | NA | NA | NA | NA | NA |
| River Salaca | 15.0 | 408.3 | 0.49 | 20.0 | 324.7 | 0.29 |
| Svartberget field infrastructure (LTER) | NA | NA | NA | NA | NA | NA |
| TERENO Wüstebach | NA | NA | NA | -2.0 | 16.7 | 0.81 |
| Uryu Experimental For | NA | NA | NA | -29.0 | 125.0 | 0.01 |
| Volbu | NA | NA | NA | NA | NA | NA |

Supplemental Figure 1. A-B. Wet deposition of ammonium and nitrate over time. Neither of these fluxes varied over time in a statistically significant way. See Figure 1 for site legend.

